

# POWDER DIFFRACTION



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## Rietveld Texture Analysis from Synchrotron Diffraction Images: II. Complex multiphase materials and diamond anvil cell experiments

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1 **Rietveld Texture Analysis from Synchrotron Diffraction Images: II. Complex**  
2 **multiphase materials and diamond anvil cell experiments**

3

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11

12 **Abstract**

13 Synchrotron X-ray diffraction images are increasingly used to characterize  
14 crystallographic preferred orientation distributions (texture) of fine-grained polyphase materials.  
15 Diffraction images can be analyzed quantitatively with the Rietveld method as implemented in  
16 the software package MAUD (Materials Analysis Using Diffraction). Here we describe the  
17 analysis procedure for diffraction images collected with high energy X-rays for a complex,  
18 multiphase shale, and for those collected *in situ* in diamond anvil cells at high pressure and  
19 anisotropic stress.

20

21 Key words: Texture analysis, Synchrotron diffraction, Rietveld method, Shale, Diamond anvil  
22 cell

23

## 24 I. INTRODUCTION

25 In a companion paper (Wenk *et al.*, 2013), we have described the basic steps for texture  
26 analysis from synchrotron diffraction images with the Rietveld method, using the software  
27 MAUD (Lutterotti *et al.*, 1997). We assume that the reader is familiar with the procedure. We  
28 will refer to some of the steps used for the general procedure from the companion paper with the  
29 corresponding number (e.g., **Part I.3**). Steps in this publication will be referred to only by  
30 number (e.g., **3**). Here we discuss complexities which arise for samples with many phases and  
31 samples which are highly deformed.

32 The first example is a sedimentary shale composed of multiple types of minerals, with  
33 different volume fractions, microstructures, and orientation distribution functions (ODFs). The  
34 second complex sample is magnesiowuestite (Mg,Fe)O measured *in situ* at ultrahigh pressure  
35 and anisotropic stress conditions with a diamond anvil cell (DAC). Keep in mind that we provide  
36 only an outline of analysis procedures. The Rietveld method and its implementation in MAUD is  
37 very general and lends itself to many applications, each of which may require slightly different  
38 approaches, modifications, and application of particular models. Data files for this tutorial can be  
39 freely downloaded from the internet (<http://eps.berkeley.edu/~wenk/TexturePage/MAUD.htm>).

## 41 II. SHALE AS AN EXAMPLE OF A COMPLEX POLYPHASE MATERIAL

### 42 A. Diffraction experiment

43 Shale is a sedimentary rock composed of many minerals. Sheet silicates comprise a large  
44 volume fraction of shales and align preferentially parallel to the bedding plane during  
45 sedimentation and compaction. Crystallographic preferred orientation (CPO) of phyllosilicates is  
46 of great interest, because it is the primary cause of elastic anisotropy observed during seismic  
47 prospecting of oil and gas deposits. Thus several studies have focused on improving synchrotron  
48 X-ray techniques to quantify textures and microstructures of shales (e.g., Wenk *et al.*, 2008;  
49 Lutterotti *et al.*, 2010; Kanitpanyacharoen *et al.*, 2011, 2012; Vasin *et al.*, 2013).

50 For this tutorial we use a sample of Kimmeridge shale from the North Sea, UK (Hornby,  
51 1998; Wenk *et al.*, 2010; Vasin *et al.*, 2013). The sample is a slab, 2mm thick (Fig. 1a). It was  
52 measured at the APS high energy beamline ID-11C during the same session as the nickel coin  
53 (Wenk *et al.*, 2013). The wavelength is 0.10798 Å, and the beam size 0.5×0.5 mm. Therefore the  
54 same instrument parameters can be applied which were obtained by refining the CeO<sub>2</sub> standard

55 (see **Part I.1-13**). During X-ray exposure, the sample was translated along the horizontal axis  
56 from  $-2.5$  to  $+2.5$  mm to increase the measured volume and rotated around the  $\omega$ -axis (Fig. 1 in  
57 Wenk *et al.*, 2013) from  $-45^\circ$  to  $+45^\circ$  in  $15^\circ$  increments (i.e., there are 7 diffraction images) to  
58 obtain adequate pole figure coverage.

59 Images were collected with a Perkin Elmer amorphous silicon detector with dimensions  
60 of  $2048 \times 2048$  pixels and a pixel size of  $200 \times 200$   $\mu\text{m}$ . The detector was approximately 1850 mm  
61 from the sample. Figure 2a shows a diffraction image with many Debye-rings from at least six  
62 major phases. Several rings display strong intensity variations due to preferred orientation.

63

## 64 **B. Analysis for axial symmetry**

65 **1. Start.** Open a new analysis in MAUD (“File  $\rightarrow$  New  $\rightarrow$  General Analysis”) and save it (“File  
66  $\rightarrow$  Save analysis as...”) in your data directory (e.g. as Shale2012-axial.par). Refining seven  
67 datasets (2D diffraction images are integrated in  $10^\circ$  steps, resulting a total of  $36 \times 7 = 252$   
68 spectra) simultaneously with a number of low-symmetry phases is time-consuming. Thus we will  
69 start with only one dataset measured at  $\omega = 0^\circ$ . The procedure with a single image is justified,  
70 because shale textures have approximately axial (cylindrical) symmetry about the bedding plane  
71 normal. By imposing this sample symmetry, complete pole figure coverage is obtained with only  
72 one dataset.

73

## 74 **2. Datasets.**

- 75 • *General:* Import calibrated diffractometer parameters (CeO<sub>2</sub>-2012.ins) as was done for  
76 the coin analysis (**Part I.15**). Restrict the refinement range to  $2\theta = 0.3 - 3.0^\circ$  (**Part I.8**).  
77 Since this shale contains several low-symmetry phases, many peaks lie at higher  $2\theta$  due  
78 to higher order reflections that strongly overlap. They do not provide much information  
79 for texture analysis, and restricting the range will greatly speed up the computation.
- 80 • *Datafiles:* Import the diffraction image (**Part I.5-7**). In ImageJ plugin, load Hornby-  
81 long-00132-0.tif, then proceed with the same steps and values as for the CeO<sub>2</sub> standard  
82 and nickel coin (**Part I.5-7**, **Part I.15**) (i.e. for image width/height size  $2048 \times 2048$ , pixel  
83 size  $0.2 \times 0.2$  mm, detector distance 1850 mm, image rotation “rotate 90 degrees left to  
84 correctly account for your sample rotation axis, etc.). Save the .esg file (e.g., as “Hornby-  
85 long-00132-0.esg”) in your data directory, and close ImageJ. As a reminder: Images need

86 to be rotated 90° left in ImageJ similarly to CeO<sub>2</sub> and nickel coin (**Part I.5**). In Figure 3  
87 (bottom) the stack of experimental diffraction spectra is displayed. The pole figure  
88 coverage is shown in Figure 1b with the pole to the bedding plane at A ( $\omega$  rotation axis).

89 • *Background function:*

- 90 ○ *Polynomial*—We will use a 4<sup>th</sup> order polynomial background common to all  
91 patterns. By default, MAUD provides three polynomial background  
92 coefficients, and two additional coefficients need to be added in the  
93 “DataFileSet” window under “Background function” tab in “Polynomial”  
94 (**Part I.8**). Right-click on each of the five background parameter values, and  
95 set them to “Refined”.
- 96 ○ *Background peaks*—It is necessary to correct for small angle scattering from  
97 platelet-shaped phyllosilicate nanoparticles, which is best visible in the  
98 diffraction image (Fig. 2a) at very low angles ( $2\theta \approx 0.1\text{-}0.2^\circ$ ) close to the  
99 beamstop. Since these platelets are oriented also small angle scattering  
100 displays azimuthal intensity variations. The broad peak extends as elevated  
101 background to the first diffraction peaks of phyllosilicates ( $2\theta \approx 0.6^\circ$ ) (Fig. 3).  
102 To correct for it, we add two symmetrical background peaks. In “Background  
103 function” go to “Background peaks” and click “Add term” (Fig. 4). The first  
104 peak is assumed to have an intensity (“Height”) of 100,000, width in  $2\theta$   
105 “HWHM” equal to  $0.1^\circ$ , width in eta (“HWHW (eta)”) of  $30^\circ$ , position along  
106  $2\theta$  (“Position”) of  $0.1^\circ$ , position along azimuth (“Position (eta)”) of  $90^\circ$ . Also  
107 set “Eta” = 1, and “Eta (eta)” = 0.5. These parameters control gaussianity of  
108 background peaks in  $2\theta$  and eta domains. Use the same parameters for the  
109 second peak, but change its “Position (eta)” to  $270^\circ$ . Double-click on  
110 background peak names to rename them. Right click on the “Height”,  
111 “HWHM” and “HWHM (eta)” boxes of both background peaks and choose  
112 the “Refined” option.

- 113 • *Save:* Make sure to save the parameter file with .par extension and in the same folder as  
114 the .esg data files.

115

116 **3. Phases.** To reduce the refinement time in this tutorial we limit our refinement to the five major  
117 phases: quartz, pyrite, kaolinite, illite-mica, and illite-smectite. There are minor phases such as  
118 feldspars with less than 5% volume and no significant texture. Quartz and pyrite structures are  
119 included in structures.mdb file of MAUD (**Part I.16**). We provide crystallographic information  
120 files (.cif) for triclinic kaolinite (Bish and Von Dreele, 1989), monoclinic illite-mica (Gualtieri,  
121 2000), and monoclinic illite-smectite (Plançon *et al.*, 1985). Refer to sections **Part I.2, I.9,** and  
122 **I.16** for information on importing .cif and working with phases. For illite-mica and illite-  
123 smectite, the first monoclinic setting has to be used (Matthies and Wenk, 2009). The provided  
124 .cif file for illite-smectite is already in the first setting. But you do need to “Edit” the illite-mica  
125 phase and under “General” change the space group from C2/c:b1 to C2/c:c1, which makes *c* the  
126 unique (2-fold) axis. Lattice parameters and atomic positions are adjusted automatically by  
127 MAUD. Rename the displays of mineral phases by double-clicking on their names.

128  
129 **4. Sample.** When multiple phases are entered, MAUD automatically assigns to each phase the  
130 same volume fraction. But with a complex sample like this shale, it is better to provide good  
131 initial estimates of phase volume fractions. Switch to “Sample” tab in the MAUD main window  
132 and “Edit” current sample. In “Layers” pick a phase from drop-down list in the right part of the  
133 window and enter a new value in the “Volume fraction” field. We start with 0.25 for quartz, 0.05  
134 for pyrite, 0.1 for kaolinite, 0.3 for illite-mica and 0.3 for illite-smectite. Weight fractions are  
135 calculated automatically by MAUD using the provided atomic structure.

136 Also, to correctly impose axial symmetry for texture, the sample orientation in the MAUD  
137 reference coordinate system must be changed. Go to “Sample position” and set the Chi value to  
138 90°. This puts the bedding plane normal of the shale into the center of the reference pole figure  
139 (B in Fig. 1b). Axial symmetry of texture in MAUD is always imposed around this axis.

140  
141 **5. Manual adjustments.** Manual adjustments of parameters are best done in “Parameter list”  
142 (“Analysis → Parameters list” or “Ctrl” and “L” on a PC or “Command” and “L” on a Mac),  
143 expanding parameters, clicking on “Value” and selecting increment to decrease or increase the  
144 value by clicking on arrows. Results are automatically updated on the “Plot” display (**Part I.12**).

145 

- Start by pressing the “Calculator” button in the MAUD main window toolbar to view  
146 calculated spectra. At the bottom of the “Plot” display there appears a list of phases with

147 each phase name of different color (unless “Black & White” is selected in “Plot options”)  
148 (Fig. 5). For each phase positions of diffraction peaks are displayed. By right-clicking on  
149 the phase name in this list, and selecting one or several phases from the menu, individual  
150 contributions of selected phases are plotted with corresponding colors.

151 • You may need to adjust some of the phase lattice parameters manually to shift position to  
152 a reasonable value. For this example, the overall calculated intensity is too low, and has  
153 to be adjusted manually to roughly match the experiment (**Part I.18**). Save your analysis  
154 before beginning refinements.

155  
156 **6. Basic refinement.** The sample is too complex to use the Wizard, and we will proceed with a  
157 mixture of semi-automatic refinement and manual (similar to **Part I.10-12**). Free and refine the  
158 parameters in the order listed below but first (in “Parameter list”) “Fix all parameters”. Then  
159 refine each parameter or parameter group until the fit has converged or is no longer visibly  
160 improving, then move on to the next parameter (usually, 4-6 iterations are enough). After you  
161 have refined a parameter, leave it free to refine, as you free and refine the next parameter in the  
162 list. If your refinement diverges, you may have to intervene with manual adjustments to bring the  
163 diverging parameters into a reasonable range. It is advised to save your analysis after each  
164 successful refinement (preferably to separate files, so that you can backtrack in case the  
165 refinement diverges and requires manual adjustments).

166 • *Intensity and backgrounds.* In “Parameter list”, click “Free scale pars”. This will set to  
167 “Refined” the overall scale of the diffraction pattern and the phase volume fractions  
168 (except the first one that is adjusted each cycle so that volume fractions add to 1). We  
169 have already set manually all five polynomial background coefficients and background  
170 peak intensity and width to “Refined” (2). Refrain from clicking the “Free backgrounds”  
171 button as this would refine all background peaks parameters, including positions that are  
172 fixed.

173 • *Thermal factors.* These diffraction data are not appropriate to refine thermal factors  
174 separately for each atom type, thus set them all equal. Click the “Bound B factors” button  
175 in the parameter list. This sets all thermal factors equal to the B factor of the first atom of  
176 the first phase in the list and sets this one to “Refined.” Run the refinement again. For  
177 these diffraction data with a narrow scattering angle range the refinement of the B-factors

178 is used as a  $2\theta$  dependent intensity adjustment, only remotely related to physical thermal  
179 vibrations.

- 180 • *Cell parameters.* All cell parameters for quartz ( $a$  and  $c$ ) and pyrite ( $a$ ) are refined since  
181 these minerals contribute several distinct sharp peaks to the diffraction pattern.  
182 Phyllosilicate peaks are weaker and strongly overlapped; thus for kaolinite and illite-mica  
183 we will not refine cell angles, but only  $a$ ,  $b$  and  $c$ . For illite-smectite, refine only the  $a$   
184 parameter. MAUD offers a user-friendly approach for the refinement of these parameters.  
185 While in “Parameters list” enter the word “cell” into a search field at the bottom. This  
186 will make only cell parameters of all phases in the sample visible in the parameters list  
187 and you may quickly set up their “Refined” status without extensive browsing.
- 188 • Make sure that you have refined the parameters specified above. Compare experiment  
189 and model on “Plot” and “Plot 2D” displays, check that all phase volume fractions are  
190 reasonable (*e.g.*, no phases have 0% volume fraction) and that cell parameters did not  
191 change significantly from their initial values.
- 192 • *Detector position.* Because the shale sample is not in the exact same position as the  $\text{CeO}_2$   
193 calibrant, the detector orientation relative to the sample is slightly different and should be  
194 refined. Set “Center displacement x,” “Center displacement y,” “Tilting error x,” and  
195 “Tilting error y” (in “Datasets → Diffraction Instrument → Flat Image Transmission”) to  
196 refine. Close all windows and run the refinement. When the refinement is done, set these  
197 four parameters back to “Fixed.”

198

199 **7. Texture and microstructure refinement.** In the “Plot 2D” display (Fig. 3), diffraction peaks of  
200 kaolinite, illite-mica, and illite-smectite show strong  $\eta$ -dependent intensity variations indicative  
201 of texture. The intensities of the quartz and pyrite diffraction peaks are almost constant, except  
202 for several increased intensity spots due to scattering from larger grains. Thus we will only refine  
203 preferred orientations of the three phyllosilicates but not of quartz and pyrite. Note that some of  
204 the peaks have incorrect width. We need to refine microstructure parameters of phases to correct  
205 this. To setup the texture and microstructure refinement, proceed with the following steps.

- 206 • Select kaolinite, edit the phase and in the “Microstructure” tab, ensure that the “Delf” line  
207 broadening model and the “Isotropic” size strain models are selected (see section VII in  
208 Wenk *et al.*, 2013, for details and references). Click the “Options” button next to “Size-



- 209 Strain model/Isotropic,” and in the new window, set both parameters (crystallite size and  
210 r.m.s. microstrain) to “Refined”.
- 211 • Go to “Advanced models”, and select the E-WIMV model for texture (**Part I.16**). Click  
212 on the corresponding “Options” button. In the “E-WIMV options panel” select 10  
213 iterations and a  $10^\circ$  for “ODF resolution in degrees”. For “Generate symmetry” select  
214 “cylindrical”. Phyllosilicates in shale display almost perfect fiber texture, thus imposing  
215 axial symmetry is justified. To speed up the refinement, click “EWIMV advanced  
216 options” button. In “Min reflex intensity” enter 0.01. It means that all the diffraction  
217 peaks with intensity below 1% of maximum peak intensity for this phase will not be used  
218 for texture calculation. Close all windows (“OK”) to save changes and return to MAUD  
219 main window. Repeat this procedure for illite-mica.
  - 220 • For illite-smectite, with a highly disordered structure, we will use the “Standard  
221 Functions” for texture model and an appropriate stacking disorder model. The so-called  
222 “Ufer single layer” model uses a super-cell approximation (Ufer *et al.*, 2004), and  
223 consequently the number of peaks fitting the pattern is increased by at least one order of  
224 magnitude. Due to an increased number of peaks and peak overlaps, the E-WIMV  
225 algorithm will run very slowly, and errors are possible when resolving diffraction peak  
226 overlaps. Thus the standard functions texture model is more appropriate in this case.
  - 227 • “Edit” the illite-smectite phase, and proceed to the “Advanced models” tab then select  
228 “Standard functions” as a texture model. Click the “Options” button. In window “Texture  
229 options panel” click the “Add term” button in the “Fiber components” panel, and set  
230 “ODF background” = 0, “Intensity” = 0.7, “ThetaY” = 0, “PhiY” = 0 (fiber axis is normal  
231 to the conventional pole figure projection plane in MAUD), “ThetaH” =  $90^\circ$ , “PhiH” = 0  
232 (fiber axis is parallel to *a* translation), “FWHM” = 30, and “Gaussian” = 0.5. Set  
233 “Intensity”, “FWHM” and “Gaussian” values to “Refined”, then click “OK”.
  - 234 • Go to the “Microstructure” tab of the phase options window. Click the “Options” button  
235 next to “Size-Strain model/Isotropic,” and in the new window change crystallite size to  
236 200 and r.m.s. microstrain to 0.01. Set both to “Refined”. Then select “Ufer single layer”  
237 as the “Planar defects model” and click “Options” button. In window “Single layer model  
238 options panel” set “Number of layers” = 10, “Stacking direction” = *a*, “Crystallite factor”  
239 = 1 (< 1 will produce additional peak broadening to (h00) type peaks), and “Microstrain

240 factor" = 1 (> 1 will produce additional peak broadening to (h00) type peaks). Refine  
241 "Crystallite factor" and "Microstrain factor".

- 242 • Leave crystallite size and r.m.s. microstrain at the default values (1000 Å and  $8E10^{-4}$ ,  
243 correspondingly) for quartz and pyrite and set these to "Refined".
- 244 • Run the refinement. At the end of the refinement, the experimental and the calculated  
245 spectra should agree fairly well in the "Plot" (Fig. 5) and "Plot 2D" (Fig. 3) displays.  
246 There are deviations in relative intensities in the "Plot" display, because this is simply an  
247 average over all spectra, not considering the relative significance of the orientation  
248 distribution. Figure 5 compares two individual spectra (obtained in "Datasets → Datafiles  
249 → View"), one with scattering vectors parallel and one perpendicular to the bedding plane  
250 normal and a good agreement for both is observed. Table 1 lists refined volume and  
251 weight fractions for the phases.

252

253 **8. Plot pole figures.** Go to the "Graphic → Texture plot" menu to view "Reconstructed intensity"  
254 of the pole figures (**Part I.20**). Select the "Active" boxes next to (001) for kaolinite (phase is  
255 chosen from "Phase" drop-down menu), and (200) for illite-mica and illite-smectite. Since we  
256 have imposed fiber symmetry, the pole figures are circular around the center with the maximum  
257 for the basal plane poles in the center (normal to bedding plane). In the "E-WIMV" window we  
258 have exported ODF's of kaolinite and illite-mica in BEARTEX format (Wenk *et al.*, 1998), then  
259 smoothed in BEARTEX with a  $7.5^\circ$  Gaussian filter and calculated pole figures. In the "Standard  
260 function" window we have also exported pole figures for illite-smectite. The pole figures are  
261 shown in Figure 6a in equal area projection. Note that illite-mica has the sharpest texture and  
262 illite-smectite shows the broadest distribution. Some data are summarized in Table 2. Save the  
263 parameter file.

264

### 265 C. Analysis without imposing texture symmetry

266 **9.** As a last step, we add the other six diffraction images to the analysis. We save the refinement  
267 which we just did under a different name (*e.g.* as Shale2012-nosymm.par).

- 268 • We create a new dataset by duplicating the available set ("Edit → Duplicate object") and  
269 "Edit" the copy by removing in "Datafiles" all datafiles in the list (select them all and  
270 click "Remove"), and then duplicate this new 'empty' dataset five more times.

- 271 • Load corresponding datafiles into the six new datasets following the procedure in **Part**  
272 **I.5-8**. Change the corresponding omega angles for each image (in “Datafiles → Modify  
273 Angles”, **Part I.15**). Rename the new datasets in the main MAUD window (*e.g.*, assign  
274 corresponding omega angle values as their names). The pole figure coverage is now as  
275 shown in Figure 1c, but considering the 90° Chi rotation (**4**), is shown in Figure 1d. Save  
276 the parameter file.
- 277 • Compute the model spectra, and review them in the “Plot” and “Plot 2D” displays. The  
278 calculated and experimental diffraction spectra should agree fairly well. Since all  
279 necessary phase parameters already have good starting values and are set to refine, there  
280 are only a few things that we need to adjust before final refinement.
- 281 • Go to “Parameter list” (“Analysis→Parameters list” menu) and click “Free scale pars” to  
282 refine incident intensity values for datasets we just added. Second, change the texture  
283 function of illite-smectite from “Standard Function” to “E-WIMV”. For each of the  
284 phyllosilicate phases, go to “E-WIMV options panel” (see **7** and **Part I.8**) and change  
285 option under “Generate symmetry” to “none”. We now have enough data to proceed  
286 without ODF symmetry imposed.
- 287 • Refine your data.

288 Pole figures of phyllosilicates, corresponding to those in Figures 6a but without imposing  
289 symmetry, are shown in Figure 6b. The pole figure coverage (Fig. 1d) indicates that during data  
290 collection the sample has been rotated around the center. The new pole figures show deviations  
291 from axial symmetry, particularly an elongation of the pole figure maximum in the vertical  
292 direction for (001) kaolinite as well as (100) illite-mica and illite-smectite. Comparing this with  
293 the coverage, we note that this distortion extends into the blind region and may be an artifact.  
294 This is further supported by the fact that maximum pole densities are higher if axial symmetry is  
295 imposed (Table 2). Only additional measurements with rotations around the other sample axes  
296 could verify if the preferred orientation pattern has perfect axial symmetry. In Figure 7 we also  
297 show pole figures (100) of kaolinite and (010) of illite-mica and illite-smectite that display a  
298 peripheral circle. Phase volume fractions for Kimmeridge shale without imposing sample  
299 symmetry are compared in Table 1 with results for axial symmetry. They are very similar.

300

301 **11.** For the Kimmeridge shale the final  $R_w$  factor is  $\approx 15\%$  and  $R_b$  is  $\approx 10\%$ . A few peaks are  
302 missing from the calculated diffraction pattern, some are too intense, and some have wrong  
303 shapes (*e.g.*, Figs. 3, 5). The missing peaks are mostly due to feldspar that could be entered into  
304 the refinement. Anisotropic crystallite shapes and microstrains could also be imposed for  
305 phyllosilicates (**Part I.7**). We have used  $\text{CeO}_2$  powder to refine instrumental parameters (Wenk  
306 *et al.*, 2013), but cerium oxide has no diffraction peaks at  $2\theta < 2^\circ$ . Thus the function describing  
307 instrumental part of diffraction peak broadening (especially the asymmetry) is poorly suited for  
308 this shale with diffraction peaks down to  $2\theta \approx 0.5^\circ$ . Parts of the instrumental peak shape function  
309 (Caglioti parameters) can be refined in this example.

310

### 311 **III. DIAMOND ANVIL CELL IN RADIAL DIFFRACTION GEOMETRY**

#### 312 **A. Experiment**

313 Rietveld texture analysis of synchrotron diffraction images can be applied to study *in situ*  
314 deformation at high pressures with a diamond anvil cell in radial diffraction geometry (rDAC)  
315 (*e.g.*, Wenk *et al.*, 2006). This proves to be an important method to determine deformation  
316 mechanisms at ultrahigh pressures, as in the deep earth (*e.g.*, Miyagi *et al.* 2010) to explain  
317 observed seismic anisotropy in the lower mantle and inner core, and to study crystal orientation  
318 changes during phase transformations (Miyagi *et al.*, 2008; Kanitpanyacharoen *et al.*, 2012b).  
319 The method can also be applied to analysis data from multi-anvil experiments such as D-DIA  
320 (*e.g.*, Wenk *et al.*, 2005).

321 The geometry of a typical rDAC deformation experiment is shown in Figure 8a,b.  
322 Diamonds not only impose pressure but also differential stress that deforms crystals in the  
323 aggregate. The diamond cell is set up in radial rather than axial geometry, i.e. the X-ray beam  
324 passes through the sample perpendicular to the compression direction so that the diffraction  
325 image records reflections from lattice planes oriented parallel and perpendicular to compression  
326 (Fig. 2c). Preferred orientation is expressed in the azimuthal intensity variations, similar to the  
327 images of the shale (Fig. 2a).

328 rDAC experiments have been performed at room temperature to pressures as high as 200  
329 GPa on iron (Wenk *et al.*, 2000) and 185 GPa on  $\text{MgSiO}_3$  post-perovskite (Miyagi *et al.*, 2010).  
330 More recently texture measurements have been made in the rDAC on magnesiowuestite

331 (Mg,Fe)O at 2273 K and  $\approx 65$  GPa, using a combination of resistive and laser heating (Miyagi *et*  
332 *al.*, 2013).

333 Contrary to the coin and shale experiments, we must take into account changes with  
334 pressure, and particularly the macroscopic stress field which imposes anisotropic elastic  
335 distortions of the lattice. As an example we use magnesiowuestite which has been previously  
336 investigated with rDAC experiments (*e.g.*, Merkel *et al.* 2002; Kunz *et al.*, 2007; Lin *et al.*, 2009;  
337 Kaercher *et al.*, 2012). This particular sample (Mg<sub>0.9</sub>Fe<sub>0.1</sub>)O has been described by Marquardt *et*  
338 *al.* (2009) and Miyagi *et al.* (2013).

339 The rDAC experiment was performed at the high pressure beamline 12.2.2. of the  
340 Advanced Light Source in Berkeley. Magnesiowuestite powder was loaded into a boron-kapton  
341 gasket. The initial sample diameter was 80  $\mu\text{m}$  with a starting thickness of 50  $\mu\text{m}$ . The sample  
342 was compressed in an rDAC, using diamond anvils with 300  $\mu\text{m}$  diameter culets (Fig. 8c).  
343 Diffraction images were recorded with a Mar3450 image plate detector, with dimensions of  
344 3450 $\times$ 3450 and a pixel size of 100 $\times$ 100  $\mu\text{m}$ , positioned approximately 285 mm from the sample  
345 and an X-ray wavelength of 0.49594  $\text{\AA}$ .

346 There are two immediate complications. First, the beam passes not just through the  
347 sample but also through a gasket, which is needed to maintain pressure. Thus there are additional  
348 diffraction lines from the gasket material, especially at low angles (Fig. 2c). Gaskets for radial  
349 DAC experiments must be made of materials that scatter as little as possible. At lower pressures,  
350 amorphous boron ( $< 100$  GPa) has been used, while at higher pressure, cubic boron nitride or  
351 beryllium have been used. For beryllium which scatters more, it is advantageous to tilt the cell  
352 slightly to have minimum beam interference. Bright diffraction spots from the diamond may also  
353 appear in the diffraction pattern. In fact, the large spot on the right side of Fig. 2c is attributed to  
354 diamond. This effect can be minimized by slightly rotating or tilting the DAC. If the cell is tilted  
355 significantly, the tilt needs to be accounted for in the analysis (Merkel, 2006). This is done by  
356 entering the appropriate sample rotation angles in MAUD.

357 A second complication is imposed anisotropic elastic strain. Lattice plane spacings are  
358 smaller in the compression direction and larger perpendicular to the compression axis. Thus, the  
359 Debye rings are not circles but ellipses. The resulting sinusoidal variations of diffraction peak  
360 positions with azimuthal angle are best seen in unrolled images (Fig. 9a, bottom).

361

## 362 B. Initial setup

363 **12. Instrument calibration.** Before analyzing the MgFeO diffraction pattern, instrument  
364 parameters should be refined with a standard sample. In this case LaB<sub>6</sub> is used  
365 (Lab6\_001.mar3450). Start a new general analysis file, and save it as “LaB6-2012.par”. The  
366 diffraction image (Fig. 2b) is quite spotty, indicative of large grain size which is not ideal,  
367 especially for refining Caglioti parameters. Before using Mar images in MAUD, they have to be  
368 converted to 16 bit TIFF images. This can be done in Fit2D (Hammersley, 1998; be aware that  
369 sometimes coordinate axis could be inverted or intensities and image size adjusted when doing  
370 image conversion with Fit2D) or preferably with the marcvr utility distributed by MarResearch.

371 Latest versions of marcvr for Linux, MacOS and Windows are freely available at  
372 <http://www.marresearch.com/download.html>. Running marcvr in Microsoft Windows systems  
373 requires Cygwin (<http://www.cygwin.com/>) with libjpeg7 and libpng12 packages. During  
374 installation of Cygwin, when prompted to select packages for the installation, uncheck “Hide  
375 obsolete packages” box. Afterwards, search for libpng12 (in All/\_obsolete subfolder) and  
376 libjpeg7 (in All/Graphics subfolder). After setting libpng12 and libjpeg7 for installation, press  
377 “Next” button and Cygwin will be installed on your computer. Put the marcvr.exe program into  
378 the “bin” subfolder of the Cygwin installation folder. You must also have Administrator rights  
379 when opening the Cygwin console and running the marcvr application.

380 After you made the appropriate .tiff image, proceed with instrumental setup in MAUD  
381 (see **Part I.1,4-13** for details). Follow directions in **Part I.1** but enter 0.49594 Å for wavelength  
382 and 285 mm as sample-detector distance. Follow step **Part I.5** and load the LaB6 TIFF image. In  
383 ImageJ specify correct Mar3450 image plate detector size (3450×3450 pixels) and pixel size  
384 (0.1×0.1 mm). Also, do not apply the image rotation (“Image→Rotate→Rotate 90 Degrees  
385 Left”) in ImageJ. We have only one diffraction image of magnesiowuestite, thus we do not need  
386 the image reorientation that was necessary to correctly account for the orientation of the  
387 horizontal sample rotation axis for beamline 11-ID-C, described for nickel coin (**Part I.5, Part  
388 I.15**) and shale (2).

389 Proceed with with **Part I.6**. The image is not well-centered, and you must manually  
390 specify estimated detector center coordinates to avoid major distortions of peak positions on  
391 different integrated spectra. You will find a reasonable alignment by setting “Center X (mm)” =  
392 172.9 and “Center Y (mm)” = 154.9. Integrate image in 5° sectors (“Number of spectra” = 72). A

393 reasonable  $2\theta$  range for the refinement is  $6.0\text{-}23.5^\circ$  For “Background function” we use  
394 “Polynomial” with five parameters (**Part I.8**).

395 Add a  $\text{LaB}_6$  phase as described in **Part I.9**. The NIST-recommended cell parameter for  
396  $\text{LaB}_6$  is  $a = 4.15689 \text{ \AA}$ , crystallite size is  $7000 \text{ \AA}$ , and microstrains are close to zero (set to  $1.0\text{E-}$   
397  $5$ ).

398 When you calculate spectra with these parameters, you will notice in the “Plot” display of  
399 the MAUD main window additional peaks at  $2\theta \approx 8.5^\circ, 13.5^\circ, 17.5^\circ$  and  $20.7^\circ$  due to some  
400 sample contamination. “Edit” current dataset, go to the “Excluded regions” tab and click “add  
401 term” button. One excluded region will appear in the list. Input values for “Min in data units”  
402 and “Max in data units” ( $2\theta$ ). We recommend excluding regions  $2\theta \approx 7.5\text{-}9^\circ, 10.5\text{-}14.5^\circ, 17.2\text{-}$   
403  $19^\circ, 20\text{-}21.3^\circ$ . This is an example of how to exclude “undesirable” data from the refinement.  
404 Note that it also excludes a couple of  $\text{LaB}_6$  peaks, overlapped with the contamination.

405 In the “Plot 2D” display, experimental diffraction peaks are hardly visible because of few  
406 very intense “spots” originating from diffraction from a few very large grains. To deal with this  
407 we select “arbitrary tex” as “Texture model” for  $\text{LaB}_6$  (**Part I.9**). In general it would be  
408 advisable not to use such a coarse-grained impure standard (Fig. 2b).

409 For the refinement of instrument parameters proceed like it is described in **Part I.11-12**.  
410 In this case you should not refine “Asymmetry” parameters as the measured diffraction peaks are  
411 far from image center and thus do not show any broadening asymmetry. Also in this case there is  
412 no eta angle dependent broadening. Repeat the refinement until convergence. Save instrument  
413 parameters as  $\text{LaB}_6\text{-2012.ins}$  (**Part I.13**).

414  
415 **13.** Start analysis for the magnesiowuestite DAC image. First, import the instrument parameters  
416 which you have just refined ( $\text{LaB}_6\text{-2012.ins}$ ). Next, convert the diffraction image of  
417 magnesiowuestite ( $\text{MgFeO\_25keV\_Z3\_004.mar3450}$ ) into .tiff format (**11**), and import it into  
418 MAUD following the same procedure as you just did for  $\text{LaB}_6$ . Unlike the coin and the shale  
419 samples, here you have only one diffraction image. Integrate it in  $5^\circ$  sectors to make 72 .esg files  
420 with individual spectra (in ImageJ). This smaller integration step is essential in this case, because  
421 there is a sharp texture and significant peak shifts due to anisotropic stress. If the integration step  
422 is too large, the variations of diffraction peak positions and intensities can not be accounted for  
423 properly. Note when the red (tracker) circle is superposed in ImageJ that the diffraction rings are

424 not circles but ellipses. They are elongated in the vertical direction (compression direction in the  
425 DAC). Also, the sides of the round image are cut by straight shadows originating from the rDAC  
426 cell. Do not forget to select the same image center as for the LaB<sub>6</sub> standard when in ImageJ  
427 (“Center X (mm)” = 172.9 and “Center Y (mm)” = 154.9). This will keep your previous beam  
428 center and detector tilt calibration valid.

429 In the “General” tab in the Datasets window, choose a “Computation range” from 11° to  
430 22° (**Part I.9**) in order to include the three prominent diffraction peaks (111), (200) and (220) of  
431 magnesiowuestite (Fig. 9, 10).

432  
433 **14.** View the spectra in “Plot” and “Plot 2D” displays. In the “Plot” display the  
434 magnesiowuestite peak at  $\approx 14.3$  Å is strangely distorted and in the “Plot 2D” display, it is  
435 evident that this distortion is due to one excessively intensive peak in diffraction spectra #64 and  
436 #65. This is the diffraction peak from the diamond anvil (Fig. 2b). Remove this spectrum (“Edit”  
437 the dataset, and in “Datafiles”, select spectrum #64 (this number is indicated in brackets after the  
438 datafile name) and click “Remove”. Do the same for the spectrum #65. Alternatively these files  
439 could be kept in the analysis but not used in the refinement by unchecking the “Enabled” box in  
440 the same window. Some spectra may be much weaker (e.g., they are shielded by the beamstop).  
441 These could also be removed from the analysis or disabled.

442  
443 **15.** Setting up the background in rDAC experiments can be difficult due to scattering and  
444 absorption from gaskets and DAC absorption effects (Fig. 2c, 9, 10). In this case it is best to use  
445 interpolated backgrounds. This is done by editing the “Dataset” and going to “Background  
446 function”. Remove all parameters for the “Polynomial” function (select parameter in the panel  
447 and click “remove parameter”), then go to the “Interpolated” tab. Check the “Interpolated  
448 background” option and enter 10 in the “Points for interpolation” field. Next click “Set  
449 interpolation points manually”. A diffraction spectrum will appear in a new window. Select  
450 “Tools → Edit interpolated background points”. A set of 10 automatically created points will be  
451 displayed. Some points lie in between peaks, corresponding to the “true” background, but some  
452 are located on diffraction peaks. Right-click on a position on the spectrum and select “add” or  
453 “remove” from the drop-down menu, in order to add or remove interpolated background points.  
454 All background points should be added where intensity is minimal (Fig. 10). Warning: in some



455 Microsoft Windows and OSX systems and computers with nVidia graphic cards, display of the  
456 points may not work correctly. Close the window and reopen it to see the interpolated  
457 background points. You may need to repeat the interpolated background setup several times,  
458 until all the points are well located.

459

460 **16.** Next load the appropriate phase using a provided structure file (MgFeO.cif). MgFeO.cif is a  
461 magnesiowuestite with 10% Fe substituting for Mg, which is specified in atom site occupancies.  
462 Calculate model spectra with the “Calculator” button on the toolbar and review the experimental  
463 and calculated diffraction patterns (Fig. 9a). Due to the high pressure applied, experimental  
464 diffraction peaks are shifted toward higher  $2\theta$  values.

465

466 **17. Sample.** With only one image and three diffraction peaks, the coverage is largely insufficient  
467 to refine the ODF. But in this DAC experiment texture should have axial symmetry around the  
468 compression direction. In MAUD, imposing axial symmetry of strain and texture requires the  
469 symmetry axis to be perpendicular to MAUD reference pole figure plane. Edit “Sample” in the  
470 MAUD main window and in “Sample position” tab set the Chi value to  $90^\circ$ . This will bring the  
471 compression direction into the center of the MAUD reference pole figure. The coverage (after  
472 this rotation) is shown in Figure 8d. First the compression axis is at A and after the Chi rotation it  
473 is at B.

474

475 **18.** Once the instrument, data files, and phase information are loaded, you will need to manually  
476 adjust several parameters to get a closer fit to the data before MAUD can run refinements.

477 • *Intensity.* Adjust the incident intensity (“\_pd\_proc\_intensity\_incident” parameter) until it  
478 is close to the data (**Part I.12** and **I.18**).

479 • *Backgrounds.* If using polynomial backgrounds, adjust the first global background  
480 coefficient: “\_riet\_par\_background\_pol0” parameter (located in the parameters list just  
481 above the intensity parameter) until the background intensity is close. If you are using  
482 interpolated backgrounds like suggested in (**15**), skip this step.

483 • *Unit cell parameters.* Since these experiments are at high pressure, you need to decrease  
484 the cell parameter “\_cell\_length\_a” (sections **Part I.12** and **I.17**) until the calculated peak  
485 positions roughly correspond to observed peaks.

- 486 • *Crystallite size and isotropic microstrain.* Adjust the crystallite size  
 487 “\_riet\_par\_cryst\_size” and microstrain “\_riet\_par\_rs\_microstrain” until peak width fits  
 488 **(Part I.9)**. Take into account that in the “Plot” display, broadening of peaks is partially  
 489 due to peak position variations. In this case good starting estimates for crystallite size and  
 490 microstrain are 300Å and 0.002, respectively.

491

492 **C. Stress models**

493 **19. Macrostress.** Lattice strain is due to the imposed anisotropic elastic stress and the elastic  
 494 properties of the crystal. It is exhibited as sinusoidal oscillations in peak position with azimuth.  
 495 An example of sinusoidal variations with azimuth can be seen in the “Plot 2D” display at the  
 496 bottom of Fig. 9a. At this point the calculated fit has no dependence of peak position on azimuth  
 497 because anisotropic stress has not yet been considered in the model. For radial DAC data with  
 498 anisotropic stress, check the “Plot 2D” display for a good fit for elastic stress rather than the  
 499 “Plot” display. The “Plot” display shows an average over all spectra, making peaks appear  
 500 broader due to angle-dependent lattice strain.

501 There are four models in MAUD that can be used to fit lattice strains, resulting in  
 502 diffraction peak shifts Two are “stress models” that convert lattice strains to macroscopic stress  
 503 tensor components, using provided elastic properties of the material. The other two models fit  
 504 lattice strain distributions and leave it up to the user to calculate stresses by hand.

505 In axial compression experiments in the DAC, the anvils impose both hydrostatic stresses  
 506 (pressure) and differential stresses. The symmetric stress tensor  $\sigma_{ij}$  can be separated into  
 507 hydrostatic  $\sigma_p$  and differential stress  $D_{ij}$  components such that:

508

$$\sigma_{ij} = \begin{bmatrix} \sigma_p & 0 & 0 \\ 0 & \sigma_p & 0 \\ 0 & 0 & \sigma_p \end{bmatrix} + \begin{bmatrix} -t/3 & 0 & 0 \\ 0 & -t/3 & 0 \\ 0 & 0 & 2t/3 \end{bmatrix} = \sigma_p + D_{ij}$$

509

510

511 where  $t$  is the axial stress component and provides lower bounds for the yield strength of the  
 512 material (Singh, 1993; Singh *et al.*, 1998). Thus, during refinement of the stresses, the  
 513 differential stresses should be constrained such that  $\sigma_{11} = \sigma_{22}$  and  $\sigma_{33} = -2\sigma_{11}$ , where  $\sigma_{33}$  is the  
 514 largest principal stress in the compression direction and is negative (corresponding to

515 compression), according to the conventions in MAUD. For the analysis described here only  
516 deviatoric stresses will be fit with the stress model. Hydrostatic stresses are accounted for by  
517 refining unit cell parameters, which in turn can be converted to pressure by utilizing an  
518 appropriate equation of state (see below). The reason for treating these separately is that  
519 differential stresses will be calculated assuming a linear stress-strain relationship which is only  
520 applicable for small strains. The volume changes of the unit cell due to pressure effects are  
521 significantly larger than those due to differential stress, and it is best to use an equation of state  
522 which properly accounts for the nonlinearity of stress-strain dependence at larger compressions.  
523 On the other hand, for the analysis of the residual stresses, *e.g.*, in engineering materials, where  
524 stress tensor components values are often within a 0.5 GPa range, it is appropriate to keep initial  
525 lattice parameters fixed. One should then only fit either stress or strain values.

526 To set up a stress model, select the phase from the “Phases” tab in MAUD main window  
527 and edit it. In “Advanced models” select one of the following models from the drop-down list in  
528 the “Strain” section: 1) “Triaxial Stress Isotropic E”, 2) “Moment Pole Stress”, 3) WSODF Popa-  
529 Balzar, or 4) “Radial Diffraction in the DAC”.

530 We will briefly describe each model and how to set it up for the refinement.

531

532 **20. Triaxial Stress Isotropic E.** This model assumes an isotropic material and as a result  
533 generally does a poor job of fitting stresses in textured material with pronounced single-crystal  
534 elastic anisotropy. But fairly isotropic materials (*e.g.*, zirconium) can be described with it and it  
535 can be applied for materials where single crystal elastic properties are unknown and you are not  
536 concerned about precise determination of stresses and want to fit textures only.

- 537
- 538 • Once you have selected “Triaxial Stress Isotropic E” click the “Options” button next to it.
  - 539 • Enter the Young’s modulus and Poisson ratio for the material being analyzed. To obtain  
540 an accurate estimate of stresses, you will need to use the Young’s modulus and Poisson  
541 ratio derived at actual high pressure value.
  - 542 • Fix the macrostresses with  $\sigma_{ij} = 0$  for  $i \neq j$ ,  $\sigma_{11} = \sigma_{22}$ , and  $\sigma_{33} = -2\sigma_{11}$ . To fix parameters,  
543 right click on the “Macrostress22” value and select “Equal to” from the drop-down menu.  
544 This will open the “Parameter binding” window. Scroll down until you find a parameter  
545 named “\_rista\_macrostress\_11” for your current phase and select it from the list (in this  
case with only one phase, it should be near the very end of the parameter list). Enter 1 in

546 the “x” box, and 0 in the “+” box. This means that “Macrostress22” will be set to  $1 * x$ ,  
 547 “Macrostress11” + 0. Set “Macrostress33” equal to (-2) times “Macrostress11”. Now that  
 548 the stress interdependencies are set, adjust or refine only “Macrostress11”. All other  
 549 stress values will be adjusted automatically.

550 • Next, manually adjust “Macrostress11” until the sinusoidal variation in peak position in  
 551 the calculated spectra is close to variations in the observed spectra in the “Plot 2D”  
 552 display. Generally the off-axis compressive stress should be less than on-axis, thus the 11  
 553 differential stress components should be positive.

554

555 **21. Moment Pole Stress.** This model is based on (Matthies *et al.*, 2001) and requires the elastic  
 556 tensor ( $C_{ij}$ ), corrected for pressure, for the material of interest (Figure 11). This is the most  
 557 sophisticated model of the four. It calculates diffraction elastic constants for each diffraction  
 558 peak of the material, taking preferred orientation into account, using different micromechanical  
 559 models similar to those used for calculating bulk polycrystal properties (*i.e.*, Voigt, Reuss, Hill,  
 560 GEO). The only difference is that for calculation of diffraction elastic constants crystal properties  
 561 should be averaged, using “moments” of ODF or pole figures (corresponding values weighted by  
 562 sine or cosine values of certain angles).

- 563 • Once you have selected the “Moment Pole Stress” model, click the “Options” button to  
 564 open the “Moment pole figures options panel”.
- 565 • Fix the macrostresses as described above with  $\sigma_{ij} = 0$  for  $i \neq j$ ,  $\sigma_{11} = \sigma_{22}$ , and  $\sigma_{33} = -2\sigma_{11}$ .  
 566 Again you will only adjust or refine “Macrostress11”.
- 567 • Enter the single crystal stiffnesses ( $C_{ij}$ , in two-index Voigt notation) for the material of  
 568 interest. Note that the elastic tensor is in general pressure-dependent, and it is important  
 569 to enter correct values for the given pressure, or the stresses obtained from the model will  
 570 be incorrect.
- 571 • In the “Stress/strain model” box in the top right choose the desired stress model. Voigt,  
 572 Reuss, Hill, PathGEO, and BulkPathGEO are available. If you choose the Hill model you  
 573 can adjust the “Weight (Voigt-Reuss)” value to get resulting arithmetically averaged  
 574 elastic constants closer to either Voigt or Reuss boundary. In almost all cases the  
 575 BulkPathGEO model is preferable. It satisfies a physically grounded “inversion relation”  
 576 (*i.e.*, stiffness is inverse compliance) for each aggregate of grains contributing to a

577 diffraction peak and also for the sample on the macroscopic scale (the latter is not true for  
578 the simpler PathGEO model).

- 579 • You can include the effects of texture by checking the “Use texture ODF” box. In fact,  
580 you are strongly advised to always do so, since differently oriented crystallites are  
581 deformed differently by the mean stress field, and their deformation in general depends  
582 on elastic constants of the material and thus on the texture. Texture effects will only be  
583 accounted for if the “E-WIMV” model is used for the texture description.
- 584 • Adjust the “Macrostress11” value until the sinusoidal variation in peak position  
585 resembles the data.

586

587 **22. WSODF Popa-Balzar.** The model, described in (Popa and Balzar, 2001), in its MAUD  
588 implementation does not provide the macrostress tensor components directly. It introduces an  
589 orientation distribution function for each strain (and consequently, stress) tensor component that  
590 is weighted according to ODF values. So, in addition to the crystal ODF, there are generally six  
591 weighted strain ODFs (WSODFs). These WSODFs are expressed as series expansions using  
592 spherical harmonics. This method is the least user-friendly, and an inexperienced user may find it  
593 very difficult to adjust all the values manually. Thus we give only a short outline of it.

- 594 • Once you have selected the “WSODF Popa-Balzar” model click the “Options” button  
595 next to it. This will open the “Harmonic strain options panel”.
- 596 • Select “fiber” as the “Sample symmetry”. You can control the length of the series  
597 expansion with the slider bar; it should be at least 4. Since this DAC experiment is  
598 believed to be centrosymmetric, all the off-diagonal strain components should be set to  
599 zero and not refined.
- 600 • After the series expansion coefficients are refined, macrostrain can be calculated by  
601 clicking the “Compute macrostrain” button (coefficients will appear in a console  
602 window). Strain distributions in the sample may be plotted as pole figures (in the MAUD  
603 main window, select “Graphic → Texture plot” and “Plot” the “Reconstructed strain”).  
604 The macrostress tensor may be calculated with a micromechanical model.

605

606 **23. Radial Diffraction in the DAC.** This model is also not a true “stress” model. While previous  
607 models can be applied to more complicated deformation geometries, “Radial Diffraction in the

608 DAC” can only be applied to axial compression. The main advantage of this model is that it  
609 allows the user to fit lattice strains for each peak separately whereas previous models imply that  
610 all the displacements of diffraction peaks correspond to one macrostress tensor, or they are  
611 restricted by crystal symmetry. The “Triaxial Stress Isotropic E” and “Moment Pole Stress”  
612 models may fail if plastic anisotropy of the material is high. In the case of magnesiowuestite  
613 some peaks exhibit much higher lattice strains than other peaks, and these two models may not  
614 be able to provide a satisfactory fit to the data.

615 The “Radial Diffraction in the DAC” model is built on the model of Singh (1993) and  
616 Singh *et al.* (1998). This model fits a  $Q(hkl)$  factor to each diffraction peak based on peak  
617 displacement and the angle to the principal stress axis.

- 618 • If you have selected the “Radial Diffraction in the DAC” model, click the “Options”  
619 button. This will open the “DAC radial diffraction options panel” (Fig. 12).
- 620 • The first two boxes are titled “Alpha” and “Beta”. “Beta” is the angle between the sample  
621 Z-axis and the maximum stress direction. “Alpha” is a rotation around Z. These two  
622 angles allow you to refine the orientation of the maximum stress axis in case it deviates  
623 from the sample Z-axis. This does happen fairly frequently at high pressures as the tips of  
624 diamond anvils deform elastically and become cupped. For now we will leave these set to  
625 zero.
- 626 • The following boxes titled “ $Q(hkl)$ ” are the  $Q(hkl)$  factor for each lattice plane given in  
627 the parentheses.  $Q(hkl)$  should be positive and generally falls in the range of 0.001 to  
628 0.007. A good initial estimate is 0.003. Enter it for the three peaks (111), (200) and (220)  
629 in the refinement range.
- 630 • Manually adjust each  $Q(hkl)$  until the lattice strains in the model are close to the observed  
631 data.

632

633 **24. Checking Sample Orientation.** If the stresses appear to be inverted, *i.e.*, “Macrostress11” or  
634  $Q(hkl)$  values are negative or the sinusoidal variations in peak positions are offset from the data  
635 by  $90^\circ$  (azimuth), then likely the sample orientation is wrong. Check that all the angles for each  
636 spectrum, dataset and whole sample correspond to experimental geometry. To check angles for  
637 individual spectra in a dataset, edit the dataset in the MAUD main window. Click on the  
638 “Datafiles” tab and ensure that angles under “Selected spectrum options” in the “Datafiles” tab

639 are set properly. To check sample orientation, edit the “Sample” in the MAUD main window,  
640 click on the “Sample position” tab where sample orientation angles “Omega”, “Chi”, and “Phi”  
641 can be adjusted. An alternative for assignment of certain angles to datafiles, or sample rotation in  
642 the “Sample” tab is to rotate the diffraction image in ImageJ when creating the .esg files, e.g. by  
643 90°. As we mentioned earlier (17), for the refinement the compression axis needs to be in the  
644 center of the pole figure.

645 You can check if the sample orientation is correct by going to “Graphic → Texture plot”  
646 in the main MAUD window. Select “Pole figure coverage” and “Plot” it. By disabling and re-  
647 enabling different spectra in the current dataset and plotting actual pole figure coverage you may  
648 see how different angles and orientations are related. In our case the pole figure coverage is  
649 shown in Figure 8d.

650

651 **25. Correcting Young’s Modulus and Poisson Ratio or  $C_{ij}$  to Pressure.** As mentioned above, you  
652 must correct the elastic moduli to pressure should you choose to use the “Triaxial Stress  
653 Isotropic E” or the “Moment Pole Stress” models. Elastic moduli are pressure dependent and  
654 often become larger as pressure increases or may display critical behavior near phase transitions.  
655 To correct elastic moduli for pressure, you will need an appropriate equation of state for your  
656 sample and a set of elastic moduli either calculated or experimentally determined for a range of  
657 pressures for your material. If your experiment is also at high temperature, you will need to  
658 correct for this as well and will also need to know the temperature dependence of elastic  
659 constants. In addition, you must account for possible anisotropic thermal expansion of the  
660 sample.

661 The easiest way to correct the elastic moduli is to create a spreadsheet which uses an  
662 equation of state, such as a 3<sup>rd</sup> order Birch-Murnaghan equation of state, to calculate pressure  
663 from the fitted unit cell parameters. Next, plot each elastic coefficient (e.g.,  $C_{11}$ ,  $C_{22}$ ,  $C_{33}$ ,  $C_{12}$  etc.  
664 or Young’s modulus and Poisson’s ratio) versus pressure. Once this is done, calculate a best fit  
665 line to each of the elastic constants and determine the equation describing the pressure  
666 dependence for each constant. This will allow you to extrapolate or interpolate elastic moduli to  
667 any reasonable pressure. Often a linear extrapolation will be good enough. Now use the pressure  
668 calculated from your unit cell parameters to determine the appropriate value of the elastic moduli  
669 using the equations for your best fit lines. You may need to perform several iterations of this

670 before the unit cell parameter and stress values stabilize. In other word you will need to  
671 calculate the pressure from the unit cell parameter, correct the elastic moduli to the pressure,  
672 input the corrected elastic moduli, and run the refinement in MAUD. After doing this you may  
673 notice that the unit cell parameter has changed. If so you will need to repeat the previous  
674 procedure until the unit cell parameter (and the corresponding pressure value) stabilizes.

675

#### 676 **D. Refinement**

677 **26. Refine Parameters.** We are ready to begin the refinement. The conditions are too complex to  
678 use the Wizard, and we will proceed with a mixture of semi-automatic refinement and manual. In  
679 the “Parameter list” click “Fix all parameters,” and save the analysis in the MAUD main  
680 window. Proceed by freeing and refining the parameters in the order listed below. Refine each  
681 parameter or parameters until the fit has converged or visibly is not getting better, then move on  
682 to the next parameter. After you have refined a parameter, leave it free to refine as you free and  
683 refine the next parameter in the list. If your refinement diverges you may have to intervene with  
684 manual adjustments.

- 685 • *Intensity and backgrounds.* Here we use interpolated backgrounds, so we refine only the  
686 intensity. In the “Parameter list” click “Free scale pars” and refine.
- 687 • *Unit cell.* In the case of cubic symmetry such as with magnesio-wuestite, we refine only  
688 “\_cell\_length\_a”.
- 689 • *Texture.* As seen in the “Plot 2D” display the texture is fairly sharp, thus we may refine  
690 the texture early. Select the “E-WIMV” texture model for MgFeO. In the “E-WIMV  
691 options panel” select 10 iterations for the texture refinement using the slider bar and a 5°  
692 resolution. For “Generate symmetry” start with “none”. Once a sample orientation and a  
693 texture that is compatible with axial symmetry is verified with pole figure plots (Fig 13a),  
694 “cylindrical” sample symmetry should be imposed (Fig. 13b). This greatly improves the  
695 effective pole figure coverage.
- 696 • *Crystallite size and r.m.s. microstrain.* Here we assume isotropic crystallite size and  
697 microstrain. In the “Parameter list” click the “Free microstructure” button and refine. As  
698 mentioned earlier, with the coarse-grained LaB<sub>6</sub> standard it is difficult to refine the  
699 instrument peak shape, and thus refined crystallite size and microstrain of



- 700 magnesiowuestite would not be accurate. Be aware that a negative crystallite size is  
701 physically meaningless and indicates that your calibration is wrong.
- 702 • *Stress model.* For “Moment pole stress” enter the elastic tensor components for your  
703 model. For magnesiowuestite at atmospheric pressure these are  $C_{11} = C_{22} = C_{33} = 279.5$   
704 GPa,  $C_{12} = C_{13} = C_{23} = 102.2$  GPa,  $C_{44} = C_{55} = C_{66} = 142$  GPa, all others are zero  
705 (Marquardt *et al.*, 2009). For “Triaxial Stress Isotropic E” our starting estimation of  
706 Young’s modulus is 290 GPa, and of Poisson ratio is 0.2. Either free “Macrostress11” if  
707 you used the “Triaxial Stress Isotropic E” or the “Moment Pole Stress” models or free the  
708  $Q(hkl)$  factors of all the peaks in the refinement range if you used the “Radial Diffraction  
709 in the DAC” model (see discussion on model selection and parameters binding for all  
710 these models in steps **20-23**). For the “Radial Diffraction in the DAC” model, some  
711 values of  $Q(hkl)$  may diverge during refinement if these peaks are of weak intensity or  
712 are not well resolved due to peak overlaps. If this happens you need to fix the appropriate  
713  $Q(hkl)$  parameter to a reasonable value, or just exclude this peak from the refinement.
  - 714 • Cell parameters and elastic constants should correspond to the same pressure. Once the  
715 stress model has converged, use the unit cell parameters to recalibrate pressure. If using  
716 the “Triaxial Stress Isotropic E” or “Moment Pole Stress” models, correct the Young’s  
717 modulus and Poisson ratio or the  $C_{ij}$ s, to the new result for pressure and refine the  
718 stresses again. Repeat this process until the stresses and unit cell parameters correspond  
719 to the same pressure. You may also do the initial refinement with “Radial Diffraction in  
720 the DAC” model. Then, once cell parameters (and thus pressure) are confirmed, you may  
721 switch to the “Moment pole stress” model and enter the pressure-corrected  $C_{ij}$ s (use Table  
722 2 from Marquardt *et al.*, 2009 to estimate pressure from the cell parameter value and their  
723 Table 3 to get  $C_{ij}$  values for this pressure). Comparison of our refinement results at this  
724 point with the equation of state (Marquardt *et al.*, 2009), suggests that our pressure value  
725 is close to 34.3 GPa (refined cell parameter is  $\approx 4.008$  Å and corresponding volume is  $\approx$   
726  $64.38$  Å<sup>3</sup>). Thus appropriate the  $C_{ij}$  values that should be used for the determination of  
727 anisotropic elastic stress tensor are  $C_{11} = C_{22} = C_{33} = 578.3$  GPa,  $C_{12} = C_{13} = C_{23} = 161.9$   
728 GPa,  $C_{44} = C_{55} = C_{66} = 141.4$  GPa, and all others are zero.
  - 729 • *Thermal factors.* Click “Bound B factors” button in the parameter list to refine them all to  
730 the same value. With our limited  $2\theta$  region we cannot refine these realistically; they will

731 serve only as an additional correction parameter to overall peak intensities. Click “Bound  
732 B factors” and refine.

- 733 • *Beam center*. If your stresses are not fitting well and you observe variations of peak  
734 position with angle eta, refine the beam center since it may have changed during DAC  
735 positioning. After refining, fix the beam center before continuing with the refinement.
- 736 • If there is evidence that the compression direction is tilted, then you may correct for this.  
737 If using the “Radial Diffraction in the DAC” model, refine the “Alpha” and “Beta” angles  
738 for a better fit. If using the “Triaxial Stress Isotropic E” or the “Moment Pole Stress”  
739 models, you may refine the sample orientation angles in “Sample” (refer to step 17) to  
740 bring the compression axis normal to the MAUD reference pole figure projection  
741 (corresponding to the sample Z axis). After refinement, fix these before continuing. To  
742 improve the fit further, you may also refine the occupation of oxygen atoms (as  
743 (Mg,Fe)O could be non-stoichiometric).

744

745 **27. Heterogeneities of strain in the DAC cell.** In the “Plot 2D” display you may observe  
746 asymmetry between the lower and the upper half of the measured spectra display, while refined  
747 spectra demonstrate perfect symmetry. This may be due to heterogeneities of the sample in the  
748 DAC, *e.g.* some grains on the periphery of the cell are subjected to lower pressures and  
749 deviatoric stress. To accommodate this, use only one half of the diffraction image. Since we  
750 imposed axial symmetry of texture and stress state, the entire diffraction pattern is not needed to  
751 derive a reasonably accurate ODF. For this case, you may remove, *e.g.*, spectra #36-71 from the  
752 analysis (refer to procedure for removing spectra in 5 and 14) and repeat your refinement.

753

754 **28.** At the end of the refinement the refined cell parameter is 4.0103 (1) Å and the corresponding  
755 volume is  $\approx 64.48 \text{ \AA}^3$ . For radial diffraction the lattice parameter represents the strain resulting  
756 from the hydrostatic (pressure) component of the stress tensor. The derived pressure is  $\approx 33.9$   
757 GPa and the elastic tensor  $C_{11} = C_{22} = C_{33} \approx 574.8 \text{ GPa}$ ,  $C_{12} = C_{13} = C_{23} \approx 161.2 \text{ GPa}$ ,  $C_{44} = C_{55} =$   
758  $C_{66} \approx 141.7 \text{ GPa}$ ; the differential macrostress 11 component is  $\approx 1.26 \text{ GPa}$ .

759 In this analysis we have been mainly concerned with preferred orientation which is  
760 displayed as pole figures (Fig. 13a,b). For materials with fiber symmetry inverse pole figures that  
761 represent the probability of the fiber axis relative to crystal coordinates is an efficient way to

762 represent textures. Figure 13c is the inverse pole figure of the compression direction plotted in  
763 MAUD and Figure 13d the inverse pole figure after processing with BEARTEX. The texture is  
764 rather sharp with a pole density maximum of  $\approx 11.5$  multiples of a random distribution, located  
765 close to 001 (Fig. 13d), as observed previously (*e.g.*, Merkel *et al.*, 2002; Lin *et al.*, 2009).

766

#### 767 **IV. CONCLUSIONS**

768         Synchrotron X-rays provide a powerful method for quantitative texture analysis of  
769 materials. Depending on sample size, beam size and wavelength, small ( $< 100 \mu\text{m}^3$ ) to large  
770 volumes ( $> 200 \text{mm}^3$ ) can be analyzed, and different sample equipment can be used to impose  
771 different conditions on the sample (*e.g.*, high pressure, high temperature, anisotropic stress).  
772 Compared to neutron diffraction, electron backscatter diffraction or pole-figure goniometry, data  
773 acquisition is fast, but data analysis is non-trivial. For complex polyphase materials (such as the  
774 shale sample) a careful manual procedure is necessary. Further complications arise for high  
775 pressure experiments, where anisotropic stresses need to be accounted for. MAUD incorporates a  
776 set of methods able to account for preferred orientations, anisotropic stresses and microstructural  
777 characteristics of material. Here we provided only a brief overview of these and simplified step-  
778 by-step procedures that give general directions for the analysis, while highlighting some possible  
779 complications. Knowledge of the instrument, sample, and experimental setup is necessary to  
780 adjust these procedures to each specific case and obtain convincing results.

781

#### 782 **Acknowledgements**

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785 from Chris Benmore and Yang Ren, as well as access to beamline 12.2.2. at ALS of Lawrence  
786 Berkeley Laboratory and help from Jason Knight. The tutorial was developed as part of a  
787 workshop at the Advanced Light Source in fall 2012, supported by CDAC, APS (BESSRC),  
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789

790

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- 876

877 **Tables**

878 **Table 1** Phase volume and weight fractions of minerals in shale (in %), with and without  
 879 imposed axial symmetry of texture.

880

	Vol. Axial	Wgt. Axial	Vol. No symm.	Wgt. No symm.
Kaolinite	6.8	6.2	7.2	6.6
Illite-mica	23.7	23.7	26.2	26.1
Illite-smectite	31.4	29.2	30.5	28.3
Quartz	31.9	29.9	29.9	27.9
Pyrite	6.2	11.0	6.3	11.1

881

882 **Table 2** Texture information for phyllosilicates in shale after processing in BEARTEX, pole  
 883 densities in m.r.d.

884

	Max axial	Min axial	Max No symm.	Min No symm
Kaolinite 001	6.84	0.22	5.14	0.31
Illite-mica 100	8.50	0.12	7.78	0.25
Illite-smectite 100	3.83	0.39	3.70	0.30

885

886 **Table 3** Texture information for magnesiowuestite after processing data in BEARTEX; pole  
 887 densities of different pole figures and inverse pole figure (IPF) in m.r.d.

888

	Max	Min		
001	5.70	0.33		
110	1.59	0.47		
111	2.11	0.12		
IPF	10.74	0.11		

889

890

891 **Figure Captions**

892

893 Figure 1. (a) Slab of shale embedded in epoxy and mounted on a pin. (b) Pole figure coverage  
894 with a single image, bedding plane normal is at A. When cylindrical symmetry is imposed  
895 bedding plane normal is rotated to B by a chi rotation; each point covers a circle around B on the  
896 pole figure (c) Coverage with seven images recorded at different sample tilts omega, (d) after 90°  
897 rotation around chi to bring the bedding plane normal into the center.

898

899 Figure 2. 2D synchrotron diffraction images. (a) Kimmeridge shale with many phases, some with  
900 strong preferred orientation. (b) LaB6 standard, rather coarse-grained and with some impurities.  
901 (c) Radial diffraction DAC experiment on magnesiowuestite. Arrow points to a diffraction spot  
902 from diamond.

903

904 Figure 3. Stack of diffraction spectra for Kimmeridge shale, omega = 0 tilt image. Experimental  
905 data at bottom and Rietveld fit on top. Some diffraction lines are labeled.

906

907 Figure 4. Window in MAUD to define background peaks.

908

909 Figure 5. Two diffraction spectra of Kimmeridge shale with scattering lattice planes parallel to  
910 bedding plane on top and perpendicular to it at bottom. Crosses are measured data and line is  
911 Rietveld fit. Below the spectra is a list of contributing phases and their corresponding diffraction  
912 peak positions are marked with ticks.

913

914 Figure 6. Pole figures of basal planes of kaolinite, illite-mica and illite-smectite for Kimmeridge  
915 shale. (a) Derived from a single image, imposing fiber symmetry. (b) Result for 7 images without  
916 imposing symmetry. The corresponding pole figure coverage is shown in Fig. 1d. Equal area  
917 projection on the bedding plane, contours in multiples of a random distribution.

918

919 Figure 7. Pole figures 100 of kaolinite and 001 of illite-mica and illite-smectite for Kimmeridge  
920 shale without imposing sample symmetry. The corresponding pole figure coverage is shown in



921 Fig. 1d. Equal area projection on the bedding plane, contours in multiples of a random  
922 distribution.

923

924 Figure 8. (a,b) Schematic sketch illustrating the geometry of deformation experiments in a  
925 diamond anvil cell in radial diffraction geometry. (c) Actual diamond culets compressing a  
926 sample contained by a gasket. (d) Pole figure coverage for the magnesiowuestite DAC  
927 experiment. A gap is visible where two spectra are disabled because they contain a diffraction  
928 spot from the diamond.

929

930 Figure 9. Measured (bottom) and calculated (top) diffraction spectra for magnesiowuestite; (a) at  
931 the beginning of the refinement. Lattice parameters are wrong and there is no texture or  
932 anisotropic stress in the model. Also note the black diffraction spot from diamond. (b) At the end  
933 of the refinement there is an excellent match in position, width and intensity. The compression  
934 direction is indicated by the arrow in (a) (larger  $2\theta$  angle corresponding to smaller d-spacing).

935

936 Figure 10. Plot display of diffraction pattern of magnesiowuestite. Dots indicate user-defined  
937 positions for background interpolation.

938

939 Figure 11. MAUD window for moment pole figures option to use as a stress/strain model.

940

941 Figure 12. MAUD radial diffraction option panel for stress-strain refinement.

942

943 Figure 13. Texture information for magnesiowuestite at 34 GPa represented as pole figures (a-b)  
944 and inverse pole figures (c-d). (a) Pole figures without imposing sample symmetry. (b) Pole  
945 figures imposing fiber symmetry. (c) Inverse pole figure of the compression direction plotted by  
946 MAUD. (d) Inverse pole figure after processing data in BEARTEX. Equal area projection,  
947 contours in multiples of a random distribution.

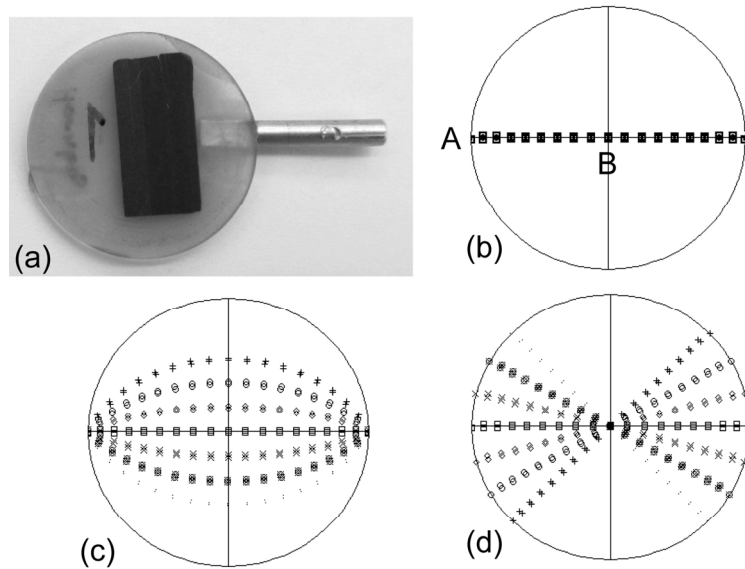


Figure 1. (a) Slab of shale embedded in epoxy and mounted on a pin. (b) Pole figure coverage with a single image, bedding plane normal is at A. When cylindrical symmetry is imposed bedding plane normal is rotated to B by a chi rotation; each point covers a circle around B on the pole figure (c) Coverage with seven images recorded at different sample tilts  $\omega$ , (d) after  $90^\circ$  rotation around chi to bring the bedding plane normal into the center.

173x107mm (300 x 300 DPI)

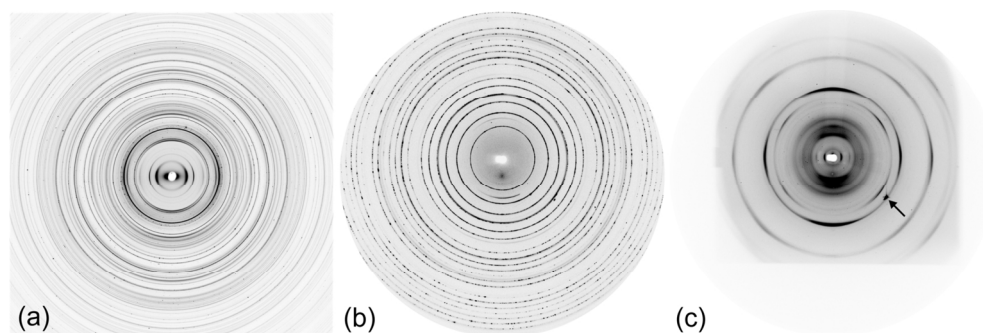


Figure 2. 2D synchrotron diffraction images. (a) Kimmeridge shale with many phases, some with strong preferred orientation. (b) LaB6 standard, rather coarse-grained and with some impurities. (c) Radial diffraction DAC experiment on magnesiowuestite. Arrow points to a diffraction spot from diamond.  
147x51mm (300 x 300 DPI)

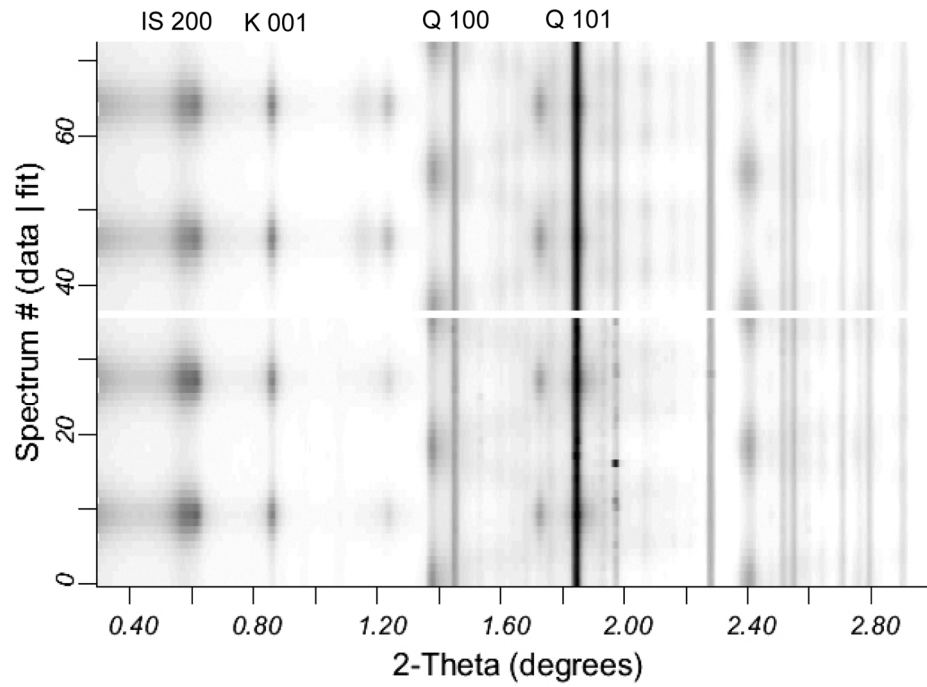


Figure 3. Stack of diffraction spectra for Kimmeridge shale, omega = 0 tilt image. Experimental data at bottom and Rietveld fit on top. Some diffraction lines are labeled.  
173x121mm (300 x 300 DPI)



Figure 4. Window in MAUD to define background peaks.  
18x12mm (300 x 300 DPI)

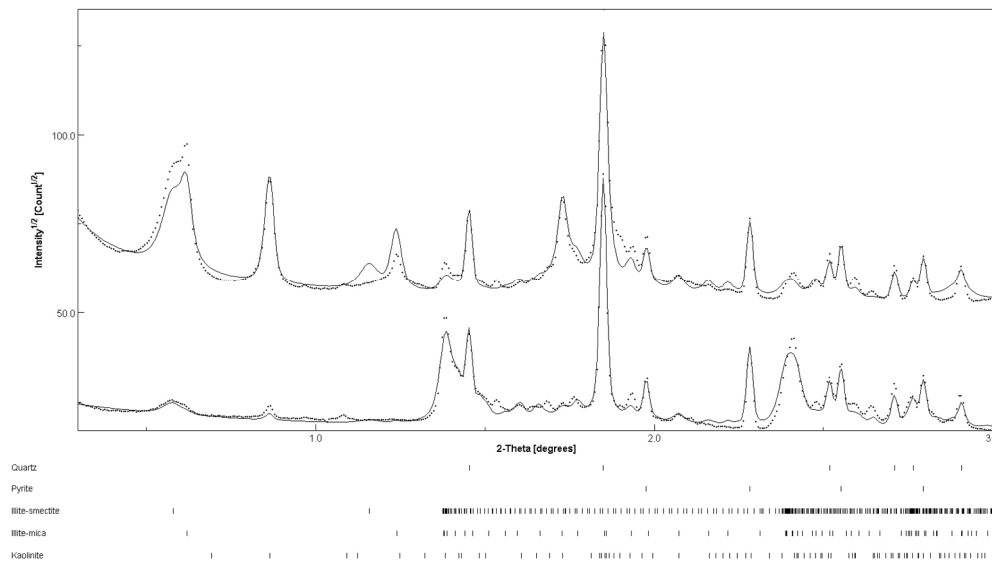


Figure 5. Two diffraction spectra of Kimmeridge shale with scattering lattice planes parallel to bedding plane on top and perpendicular to it at bottom. Crosses are measured data and line is Rietveld fit. Below the spectra is a list of contributing phases and their corresponding diffraction peak positions are marked with ticks.

213x123mm (200 x 200 DPI)

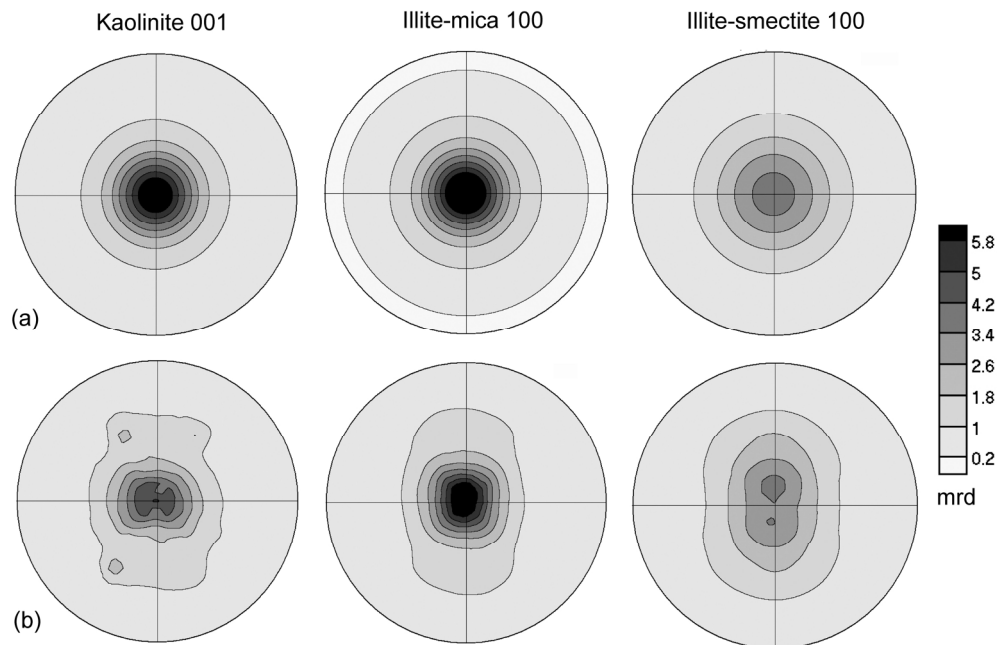


Figure 6. Pole figures of basal planes of kaolinite, illite-mica and illite-smectite for Kimmeridge shale. (a) Derived from a single image, imposing fiber symmetry. (b) Result for 7 images without imposing symmetry. The corresponding pole figure coverage is shown in Fig. 1d. Equal area projection on the bedding plane, contours in multiples of a random distribution.  
198x127mm (300 x 300 DPI)

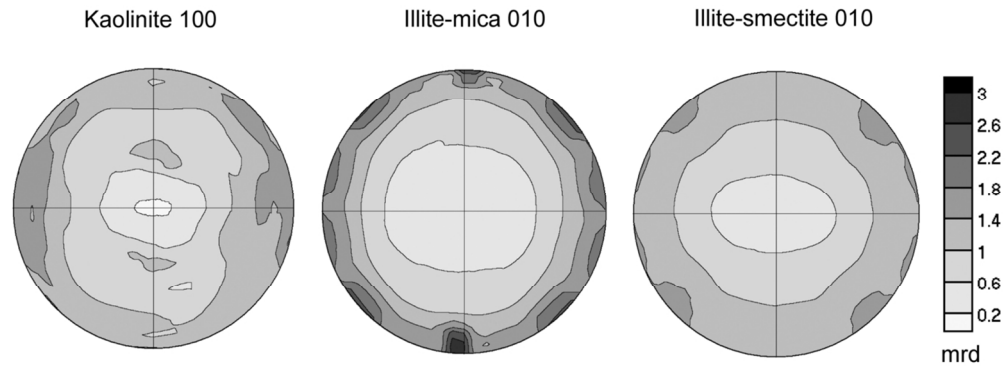


Figure 7. Pole figures 100 of kaolinite and 001 of illite-mica and illite-smectite for Kimmeridge shale without imposing sample symmetry. The corresponding pole figure coverage is shown in Fig. 1d. Equal area projection on the bedding plane, contours in multiples of a random distribution.  
114x43mm (300 x 300 DPI)



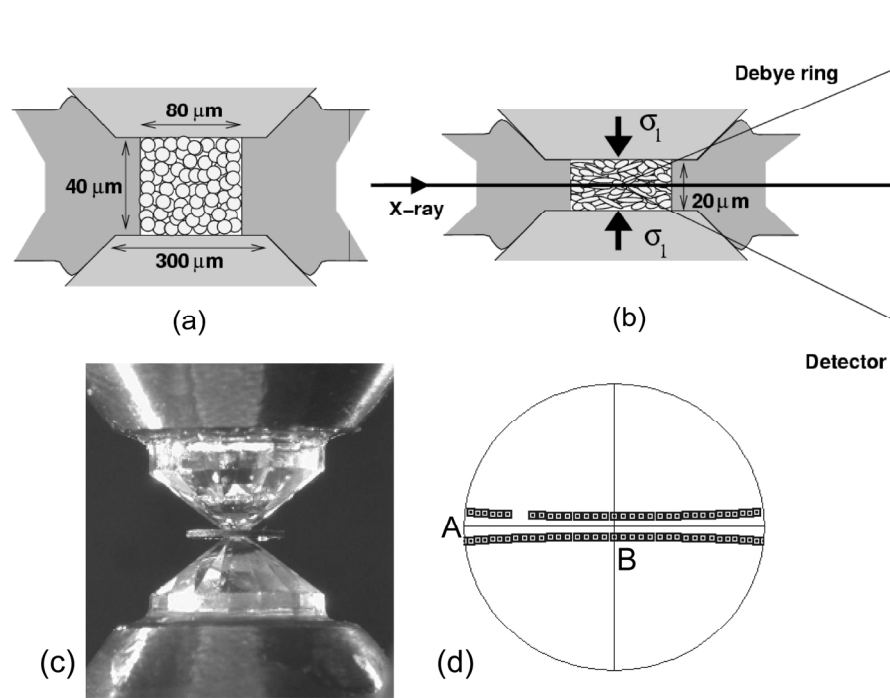


Figure 8. (a,b) Schematic sketch illustrating the geometry of deformation experiments in a diamond anvil cell in radial diffraction geometry. (c) Actual diamond culets compressing a sample contained by a gasket. (d) Pole figure coverage for the magnesiowuestite DAC experiment. A gap is visible where two spectra are disabled because they contain a diffraction spot from the diamond.

199x145mm (300 x 300 DPI)

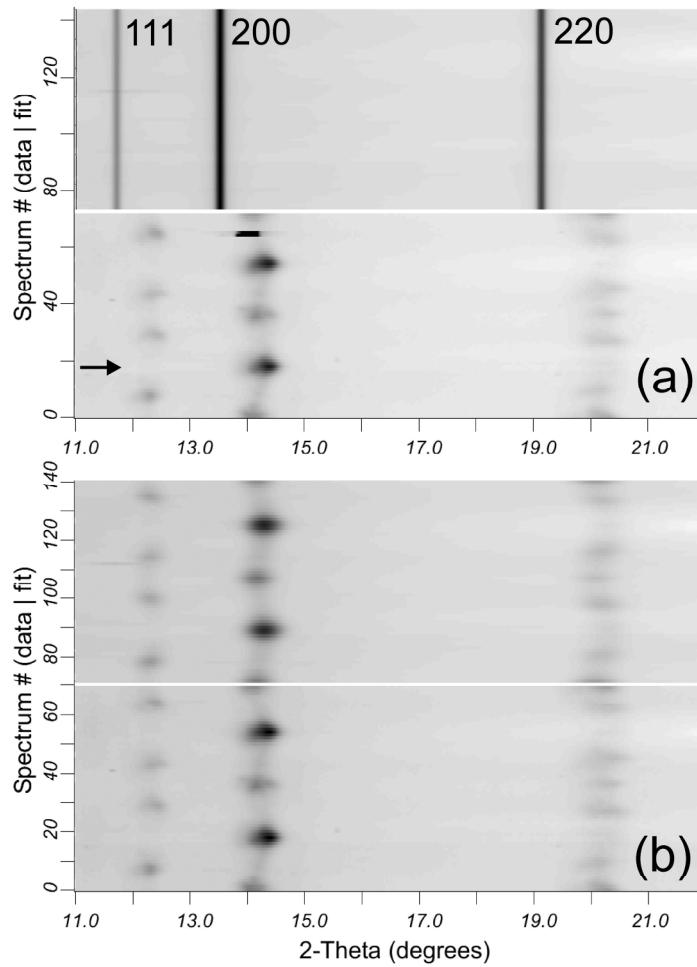


Figure 9. Measured (bottom) and calculated (top) diffraction spectra for magnesiowuestite; (a) at the beginning of the refinement. Lattice parameters are wrong and there is no texture or anisotropic stress in the model. Also note the black diffraction spot from diamond. (b) At the end of the refinement there is an excellent match in position, width and intensity. The compression direction is indicated by the arrow in (a) (larger  $2\theta$  angle corresponding to smaller d-spacing).

227x221mm (300 x 300 DPI)

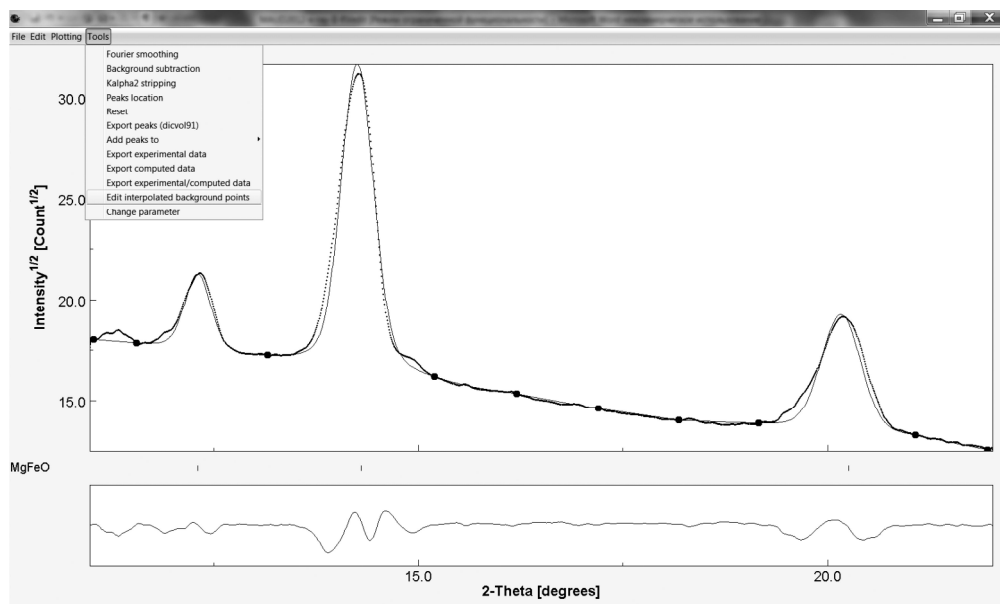


Figure 10. Plot display of diffraction pattern of magnesiowuestite. Dots indicate user-defined positions for background interpolation.  
1482x884mm (72 x 72 DPI)

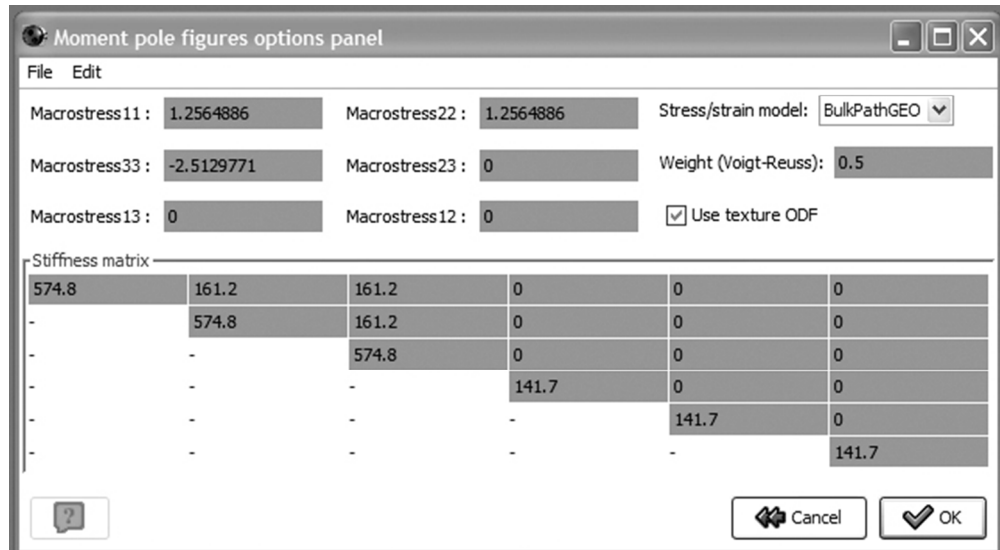


Figure 11. MAUD window for moment pole figures option to use as a stress/strain model. 227x124mm (72 x 72 DPI)

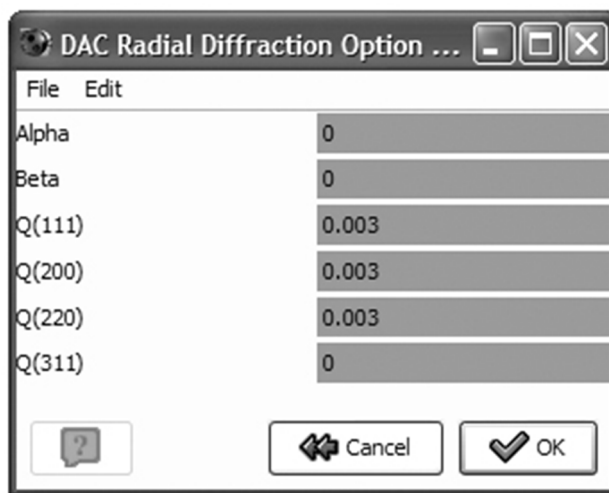


Figure 12. MAUD radial diffraction option panel for stress-strain refinement.  
110x87mm (72 x 72 DPI)

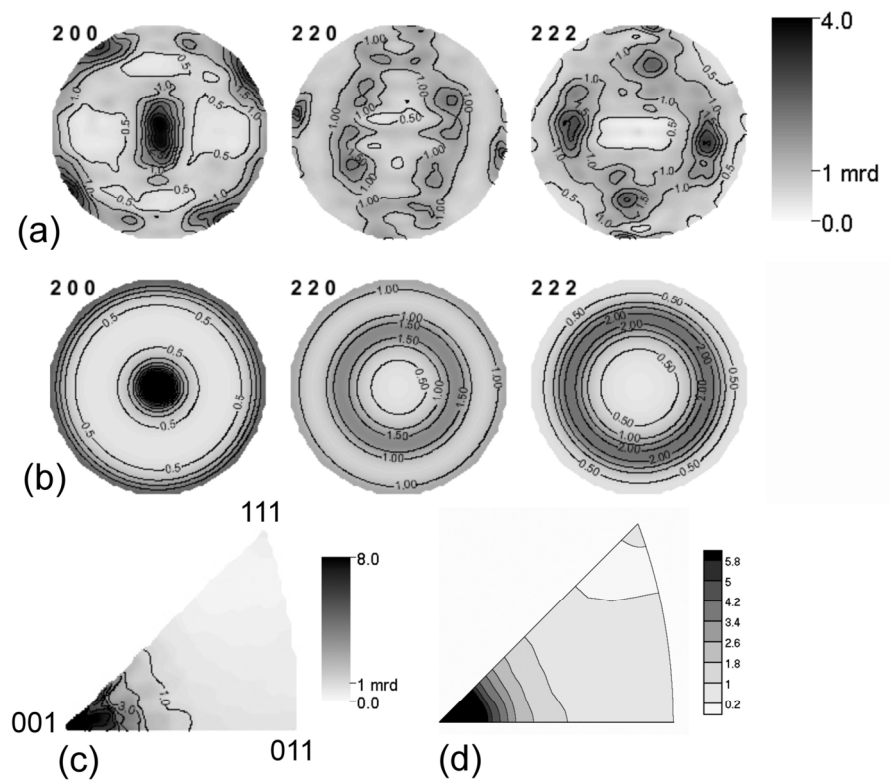


Figure 13. Texture information for magnesiowuestite at 34 GPa represented as pole figures (a-b) and inverse pole figures (c-d). (a) Pole figures without imposing sample symmetry. (b) Pole figures imposing fiber symmetry. (c) Inverse pole figure of the compression direction plotted by MAUD. (d) Inverse pole figure after processing data in BEARTEX. Equal area projection, contours in multiples of a random distribution.

157x126mm (300 x 300 DPI)