

# Microtwinning hypothesis for a more ordered vaterite model

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An orthorhombic fully ordered structural model is proposed for vaterite [space group  $Ama2$ ,  $a = 8.4721(5)$  Å,  $b = 7.1575(7)$  Å,  $c = 4.1265(4)$  Å,  $Z = 4$ , and  $V = 250.23(4)$  Å<sup>3</sup>]. It is based on a microtwinning hypothesis, with three domains rotated by 120° along the orthorhombic  $a$  axis, regenerating a pseudo-hexagonal habit. The solution came from direct space *ab initio* calculations applied to the powder diffraction data. However, five weak superstructure reflections seen in single-crystal and powder diffraction experiments, leading to a six times larger unit cell, are still unexplained. © 2011 International Centre for Diffraction Data. [DOI: 10.1154/1.3552994]

Key words: vaterite, powder diffraction, *ab initio*, direct space, twinning

## I. INTRODUCTION

Among the three crystallized polymorphs of CaCO<sub>3</sub>, vaterite is the least stable form under natural conditions, with around 3 kJ/mol difference compared to calcite, and is potentially involved in the first steps of crystallization of the two others (calcite and aragonite). Its complete structural determination would consequently shed important light to understand scaling formation and biomineralization processes. Polymorphs' formation and modification of inorganic, biogenic, and biomimetic vaterite crystal habits can be induced using special conditions or additives. *In vitro* crystallization of vaterite was achieved using *Mytilus californianus* nacre macromolecules and supersaturated solutions (Falini *et al.*, 1996). Mg<sup>2+</sup> ions (Chen *et al.*, 2005) inhibit vaterite formation while NH<sub>4</sub><sup>+</sup> ions (Gehrke *et al.*, 2005) generate hexagonal polydomain vaterite crystals. Surfactants (Dupont *et al.*, 1997) and/or polymers (Wei *et al.*, 2004) also strongly affect crystal habits even in small amounts in the solution. Calcite and aragonite biocrystals, stabilized under organic influences, were shown to exhibit unit-cell and structural distortions (Pokroy *et al.*, 2004, 2006, 2007), depending on the biocrystal habits (Ouhenia *et al.*, 2008a). The vaterite structure might also distort under thermodynamic and/or organic variations, and this would also be the case of other microstructural defects such as microtwins, recently observed in vaterite tablets of freshwater lacklustre pearls (Qiao and Feng, 2007). Trying to understand the growth of biominerals, our previous results indicated that synthesis of calcium carbonate in the presence of polyacrylic acid (PAA) can shift the chemical equilibrium between the allotropic forms, increasing vaterite and aragonite fractions depending on the temperature used. PAA can also control crystal sizes, quality, and shapes but we also observed such crystal habit variations when no PAA was added (Ouhenia *et al.*, 2008b). This led us

to a reinvestigation of the vaterite structure under purely inorganic mineral formation and for which only very old experimental atomic coordinates are available in databases (Gražulis *et al.*, 2009). In the more accepted CaCO<sub>3</sub> vaterite structural model based on single-crystal data (Kamhi, 1963), the CO<sub>3</sub><sup>2-</sup> ions are randomly distributed over three orientations parallel to the  $c$  axis ( $P6_3/mmc$  space group,  $a' = 4.13$  Å,  $c' = 8.49$  Å, and  $Z = 2$ ). However, that model did not take into account the five weak superstructure reflections, leading to a cell having a volume six times larger ( $a = a'\sqrt{3}$ ,  $c = 2c'$ , and  $Z = 12$ ). Previously, Meyer (1959) proposed tentative atomic coordinates in an orthorhombic cell with dimensions  $a_0 = 4.13$  Å,  $b_0 = 7.15$  Å, and  $c_0 = 8.48$  Å and space group  $Pbnm$ , but this was not checked against any diffraction data. From first-principles calculations and molecular dynamic simulations applied in order to determine the local order, a hexagonal superstructure ( $P6_522$ ) was recently proposed (Wang and Becker, 2009) with  $Z = 18$  and cell dimensions of  $\sqrt{3}$  times in  $a'$  and three times in  $c'$ . In this latter article, the controversy about the vaterite structure is detailed.

To the best of our knowledge, merohedral or pseudomerohedral microtwinning models have not been considered yet in order to try to explain the single-crystal data with a fully ordered model. We explore this possibility here, giving the highest priority to the fitting of the experimental powder diffraction data by models consistent with plausible interatomic distances, coordinations, and angles and compatible with the previous single-crystal observations, which are obviously prerequisites.

## II. EXPERIMENTAL AND CHECKING PREVIOUS MODELS

The vaterite sample was synthesized using a procedure similar to that of Sato and Matsuda (1969): a 3M solution of CaCl<sub>2</sub> in 50 ml of de-ionised water and 1M solution of

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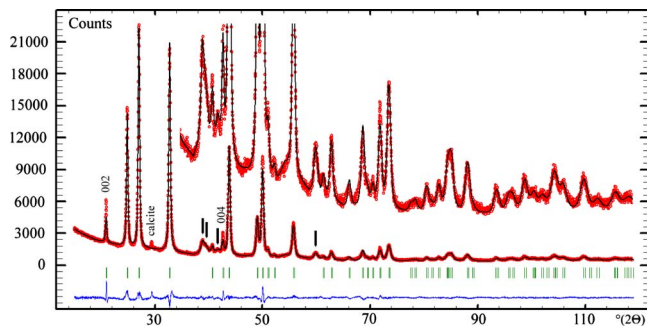


Figure 1. (Color online) Le Bail fit with the small hexagonal cell of Kamhi (1963) (space group  $P6_3/mmc$ ). There are four clear additional peaks noted by vertical arrows which are reproduced by background points (corresponding from lower to larger  $2\theta$  angles to the 211, 205, 213, and 401 reflections in the supercell). The fifth supercell reflection seen with the single-crystal data of Kamhi is closely overlapping with an intense reflection of the subcell ( $\sim 33^\circ 2\theta$ ). Calcite is present as very weak impurity; its most intense reflection is indicated. The 001 vaterite reflections are narrower than the others.

$K_2CO_3$  in 450 ml of de-ionised water, both preheated at  $33^\circ C$ , were mixed simultaneously in a beaker and stirred for 20 min at  $33^\circ C$ . The product was washed actively with water and methanol and then dried in a desiccator under vacuum in the presence of  $P_2O_5$ . The powder pattern was recorded on a PANalytical MPD PRO diffractometer in the Bragg-Brentano geometry, equipped with a X'Celerator detector ( $Cu K\alpha$ ). From the peak position analysis and indexing efforts with the MCMAILLE software (Le Bail, 2004), the small hexagonal cell of Kamhi (1963) is easily recovered. The supplementary peaks are really few and weak, as seen in Figure 1 from a structureless Le Bail fit (Le Bail, 2005) using the FULLPROF software (Rodríguez-Carvajal, 1993), where they were modelled as belonging to the background. Successful indexing in a new cell that fits these more broadened peaks is not easy. These peaks are present on some of the 13 ICDD PDF cards (ICDD, 2009) obtained by a search

on the vaterite mineral name PDF 00-033-0268 (ICDD, 2009), to give one with experimental data). They are also seen in previously reported synthesis (Dupont *et al.*, 1997). It was decided to check all the main published models against the powder diffraction data recorded. The  $CO_3$  triangles were the subject of soft constraints during refinements.

### A. Meyer (1959), orthorhombic $P$ , ordered model

The fit of Rietveld (1969) by using FULLPROF with the fixed original tentative atomic coordinates from the orthorhombic cell of Meyer (1959) (Table I) is not very satisfying ( $R_{WP}=36.3\%$  and  $R_B=25.7\%$ ). Refining the atomic coordinates improves the fit significantly ( $R_{WP}=18.5\%$  and  $R_B=7.90\%$ ). However, this cell is unable to explain all the superstructure peaks on the powder pattern and it adds calculated peaks incompatible with the Kamhi hexagonal subcell and the powder data. An orthorhombic cell which would not provide additional reflections has to be C face centered, not primitive, but would still not allow indexing of all the superstructure peaks. It should be recalled that Meyer never checked his structure hypothesis against the experimental data. Moreover, the current fit attempt seems to be the first, though his model continues to be cited as one of the possible structures for vaterite.

### B. Kamhi (1963), hexagonal subcell, disordered

A Rietveld refinement in the subcell using Kamhi's coordinates (Table I) leads to  $R_{WP}=13.3\%$  and  $R_B=4.57\%$  (Figure 2). Kamhi stated "the crystals are imperfect, and under high magnification can be seen to be aggregates of uniformly oriented hexagonal plates." This is a feature illustrated more recently in the work of Gehrke *et al.* (2005). If Kamhi's remark aimed at underlining the "imperfection" of the crystals, the vaterite aggregates in the work of Gehrke *et al.*, while exhibiting noticeable mosaicity at the TEM scale,

TABLE I. Atomic coordinates for the three main models from the Rietveld refinements.

Model 1. Meyer (1959), $Pbnm$ , $Z=4$ , $a=4.1291(5)$ Å, $b=7.1581(9)$ Å, and $c=8.4764(7)$ Å						
Atom	Wyckoff	$x$	$y$	$z$	$B_{iso}$	Occ.
Ca	$4a$	0	0	0	1.68(5)	1
C	$4c$	0.0794(17)	0.646(3)	1/4	4.3(1)	1
O1	$4c$	0.3722(9)	0.5890(9)	1/4	4.3(1)	1
O2	$8d$	-0.0552(8)	0.6678(14)	0.1219(3)	4.3(1)	1
Model 2. Kamhi (1963), $P6_3/mmc$ , $Z=2$ , $a=4.1304(2)$ Å, and $c=8.4749(5)$ Å						
Ca	$2a$	0	0	0	1.84(4)	1
C	$6h$	0.7074(5)	$2\times$	3/4	2.65(8)	1/3
O1	$12k$	0.6242(4)	$2\times$	0.6210(2)	2.65(8)	1/6
O2	$6h$	0.8793(3)	$2\times$	3/4	2.65(8)	1/3
Model 3. This work, $Ama2$ , $Z=4$ , $a=8.4721(5)$ Å, $b=7.1575(7)$ Å, and $c=4.1265(4)$ Å						
Ca	$4a$	1/2	1/2	$0^a$	1.86(3)	1
C	$4b$	1/4	0.640(3)	0.452(4)	3.12(8)	1
O1	$8c$	0.3795(2)	0.6711(16)	0.5636(17)	3.12(8)	1
O2	$4b$	1/4	0.5432(12)	0.1912(14)	3.12(8)	1

<sup>a</sup>Fixed coordinate. The CIF corresponding to model 3 was deposited at the Crystallography Open Database (www.crystallography.net) (Gražulis *et al.*, 2009) with entry number 3000002.

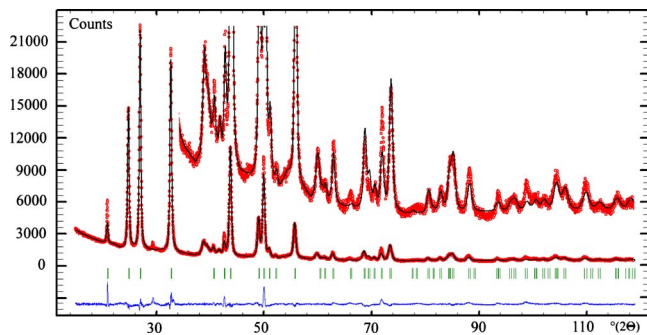


Figure 2. (Color online) Rietveld fit with the hexagonal subcell of Kamhi (1963). The superstructure peaks (see Figure 1) are modelled by background points.

still give rise to perfect single-crystal diffraction patterns. Later, Meyer (1969) reused the hexagonal supercell but the proposed coordinates do not allow constructing any reasonable model. So the Kamhi model is still considered to be the best approach to the vaterite crystal structure in spite of the considerable disorder imposed on the  $\text{CO}_3$  group.

### C. Wang and Becker (2009), $P6_522$ ordered model by MD annealing

Why propose  $3c'$  whereas Kamhi experimentally proved that only doubling  $c$  accounted for the single-crystal data? Indeed, using the  $P6_522$  model and adjusting the predicted cell parameters  $a=7.29$  and  $c=25.302$  Å to the observed ones  $a=7.16$  and  $c=25.44$  Å fails to fit the positions of the additional superstructure peaks and brings a large number of supplementary peaks not observed in any experimental powder pattern data from the literature (nor in Kamhi's single-crystal data). Rietveld refinements were not attempted in this case. Other first-principles calculations leading to models similar to the orthorhombic cell of Meyer (1959) were discarded because their cell parameters strongly disagree (up to 10%) with the experimental ones (Medeiros *et al.*, 2007).

### III. TWINNING HYPOTHESIS, STRUCTURE SOLUTION, AND RIETVELD REFINEMENTS

Would it be possible to obtain a fully ordered model which could explain both the powder pattern and the single-crystal data? Various twinning hypotheses by merohedry or pseudomerohedry were tested, targeting those not adding any supplementary reflections to the single-crystal data subcell and supercell. We stayed first in the smallest hexagonal cell ( $Z=2$ ) from Kamhi (1963) and a convincing solution was obtained by following the path of the relation group-subgroup in the hypothesis of a pseudohexagonal symmetry generated from orthorhombic C-centered cells twinned or microtwinning in three domains related by rotation of  $120^\circ$  around the  $c$  axis. The occurrence of that twinning law has been encountered many times already, in the  $\beta\text{-AlF}_3$  case (Le Bail *et al.*, 1988) for instance. The procedure consisted in extracting the  $|F_{\text{obs}}|$  from the powder pattern (Le Bail, 2005) in the various cells and space group hypothesis; then, the extracted  $|F_{\text{obs}}|$  was used for structure solution attempts in direct space by the ESPOIR software (Le Bail, 2001), moving

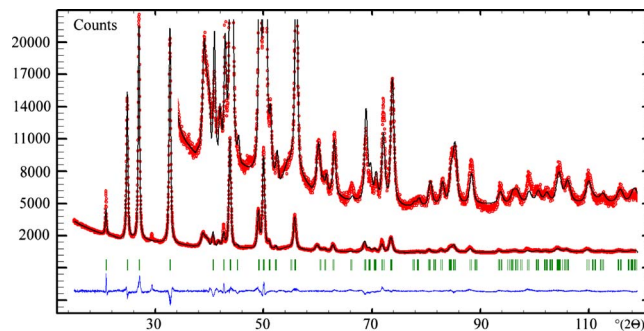


Figure 3. (Color online) The Rietveld fit with the *Ama2* fully ordered model. The superstructure peaks (see Figure 1) are modelled by background points.

independently in the cell, by a Monte Carlo process, one or several rigid  $\text{CO}_3$  groups and the Ca atoms. This twinning hypothesis is able to reconcile the Kamhi hexagonal subcell with the Meyer orthorhombic cell, provided the latter is C centered. Tests were made in  $Cmcm$  (maximum nonisomorphic subgroup for  $P6_3/mmc$ ) and  $Cmc2_1$  space groups without obtaining the expected full order, so the symmetry was finally lowered to monoclinic in the  $C2/c$  and  $Cc$  space groups. A model showing two disordered positions for the  $\text{CO}_3$  group was obtained in  $C2/c$ , and the expected full ordering was obtained in  $Cc$  with all atoms in the general Wyckoff position  $4a$ . The latter model was convincing for two main reasons: the eightfold Ca coordination satisfied usual interatomic distances and the  $\text{CO}_3$  group planes were lying all parallel (almost) to the same axis. Checking this model for missing symmetry using PLATON (Spek, 2003) finally helped to conclude that *Ama2* was the most appropriate space group, for which the Rietveld refinements resulted in  $R_{\text{WP}}=13.4\%$  and  $R_B=5.22\%$  (Figure 3). The obtained reliability factors are comparable to the ones resulting from the hexagonal subcell highly disordered model but with a fully ordered model. The refined atomic coordinates are presented in Table I together with those of the Kamhi and Meyer subcell models. Interatomic distances are in Table II and X-ray diffraction data are in Table III.

TABLE II. Selected geometric parameters (Å, deg) for the vaterite described in *Ama2*.

Ca-O2 2×	2.281(2)	O1-O2	2.098(9)
Ca-O1 <sup>a</sup> 2×	2.405(8)	O1-O1 <sup>b</sup>	2.194(2)
Ca-O1 <sup>c</sup> 2×	2.579(10)	O1-O1 <sup>d</sup>	3.189(13)
Ca-O1 2×	2.820(8)	O1-O1 <sup>e</sup>	3.115(9)
C-O1 2×	1.211(8)	O1-O2 <sup>f</sup>	2.958(9)
C-O2	1.280(19)	O1-O2 <sup>g</sup>	2.928(13)
O1-C-O1	130.0(6)	O1-C-O2	114.8(12)

<sup>a</sup>Symmetry code:  $x, y, z-1$ .

<sup>b</sup>Symmetry code:  $-x+1/2, y, z$ .

<sup>c</sup>Symmetry code:  $x, y-1/2, z-1/2$ .

<sup>d</sup>Symmetry code:  $-x+1, -y+1, z$ .

<sup>e</sup>Symmetry code:  $-x+1/2, -y+3/2, z-1/2$ .

<sup>f</sup>Symmetry code:  $x, y, z+1$ .

<sup>g</sup>Symmetry code:  $x, y+1/2, z+1/2$ .

TABLE III. X-ray diffraction data for vaterite indexed in the *Ama2* subcell (Cu  $K\alpha_1$ ).

$2\theta_{\text{obs}}$	$d_{\text{obs}}$	$I_{\text{obs}}$	$h$	$k$	$l$	$2\theta_{\text{cal}}$	$d_{\text{cal}}$	$I_{\text{cal}}$	$\Delta 2\theta$
20.953	4.2362	19.6	2	0	0	20.954	4.2360	11.5	-0.001
24.883	3.5753	62.1	0	2	0	24.859	3.5788	28.0	0.024
			0	1	1	24.886	3.5749	35.6	-0.003
27.050	3.2936	100.0	1	2	0	27.024	3.2967	54.3	0.026
			1	1	1	27.049	3.2937	45.7	0.001
32.753	2.7320	92.6	2	2	0	32.731	2.7338	18.1	0.022
			2	1	1	32.752	2.7320	84.5	0.001
38.871	2.3149	6.2			<sup>a</sup>				
39.381	2.2861	4.9			<sup>a</sup>				
40.672	2.2165	4.8	3	2	0	40.663	2.2169	0.1	0.009
			3	1	1	40.681	2.2160	7.2	-0.009
41.774	2.1605	2.9			<sup>a</sup>				
42.673	2.1171	11.2	4	0	0	42.653	2.1180	7.7	0.020
43.822	2.0642	51.3	0	3	1	43.793	2.0655	50.6	0.029
			0	0	2	43.843	2.0632	15.6	-0.021
49.049	1.8557	18.3	2	3	1	49.027	1.8565	12.7	0.022
			2	0	2	49.072	1.8549	11.9	-0.023
50.023	1.8219	47.9	4	2	0	49.997	1.8227	24.2	0.026
			4	1	1	50.012	1.8222	34.8	0.011
51.039	1.7879	4.1	0	4	0	50.995	1.7894	1.9	0.044
			0	2	2	51.054	1.7875	2.1	-0.015
52.204	1.7508	1.3	1	2	2	52.262	1.7490	1.9	-0.058
55.767	1.6470	16.1	2	4	0	55.719	1.6483	6.0	0.048
			2	2	2	55.774	1.6468	17.3	-0.007
59.898	1.5429	2.7			<sup>a</sup>				
62.842	1.4776	3.6	4	3	1	62.786	1.4787	3.8	0.056
			4	0	2	62.825	1.4779	1.5	0.017
66.111	1.4122	1.0	6	0	0	66.120	1.4120	1.5	-0.009
68.609	1.3667	3.8	4	4	0	68.601	1.3669	3.1	0.008
			4	2	2	68.650	1.3660	5.1	-0.041
68.455	1.3521	1.5	0	5	1	69.437	1.3524	0.9	0.018
			0	4	2	69.474	1.3518	0.6	-0.019
			0	1	3	69.535	1.3508	0.7	-0.080
70.533	1.3341	1.3	1	5	1	70.446	1.3355	1.1	0.087
			1	4	2	70.482	1.3349	0.7	0.051
			1	1	3	70.543	1.3339	0.4	-0.010
71.840	1.3130	5.9	6	2	0	71.810	1.3135	2.0	0.030
			6	1	1	71.822	1.3133	6.3	0.018
72.496	1.2875	7.2	2	5	1	73.435	1.2884	3.1	0.061
			2	4	2	73.461	1.2878	5.4	0.035
			2	1	3	73.531	1.2869	6.0	-0.035
80.532	1.1918	1.3	0	6	0	80.437	1.1929	1.3	0.095
			0	3	3	80.542	1.1916	1.4	-0.010

<sup>a</sup>Supercell peaks.

#### IV. DISCUSSION AND CONCLUSION

The *Ama2* model is very similar to that of Meyer, as can be seen in Figure 4. The Ca atoms are in a distorted cubic coordination in both cases. The difference lies in the positions of the C atoms for half of them in adjacent possible oxygen triangles. Two edges of the  $\text{CaO}_8$  cubes are shared by two  $\text{CO}_3$  triangles, which explain the cube distortion by short O-O distances. The structure is built of  $\text{CaO}_8$  cubes interconnected by edges like in the  $\text{CaF}_2$  fluorite structure, but expanding in only two dimensions, forming layers parallel to the *bc* planes (Figure 5) of the *Ama2* model. Each layer is connected by corners (O2 atoms) to the two neighbouring mirror-related layers (Figure 4). The  $\text{CO}_3$  groups enforce the layers connections, the carbon atom being at the center of the

triangle formed by this O2 atom and two O1 atoms from the neighbouring  $\text{CaO}_8$  cubes in two adjacent layers (Figure 6). This organization builds two neighbouring triangular possibilities for inserting the  $\text{CO}_3$  group, sharing the oxygen atom O1. If both places were occupied, then  $\text{C}_2\text{O}_5$  groups would exist which is impossible. Compared to other  $\text{CaCO}_3$  polymorphs, there is no edge sharing between the  $\text{CO}_3$  groups and the calcium  $\text{CaO}_6$  octahedra in calcite. In the latter compound, only corner sharing is observed both between octahedra themselves and between  $\text{CO}_3$  groups and octahedra. In aragonite, all three edges of the  $\text{CO}_3$  groups are common with edges of three different  $\text{CaO}_9$  polyhedra. In this sense, one can see the vaterite structure as a real intermediate form between calcite and aragonite; these two more stable poly-

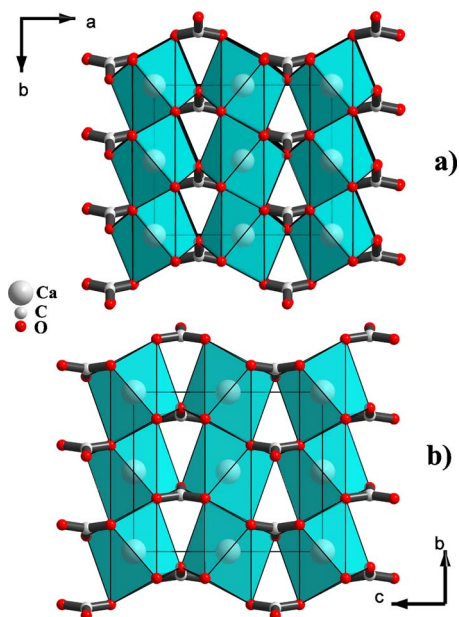


Figure 4. (Color online) Comparison of the *Ama2* model (a) with the *Pbnm* Meyer model (b). Projection along the short axis (4.13 Å) showing the layers of  $\text{CaO}_8$  cubes interconnected by edges and the  $\text{CO}_3$  groups ensuring their linkage. The  $\text{CO}_3$  groups are all pointing in the same direction by O2 in the *Ama2* model, and they alternate in opposite directions in the *Pbnm* model.

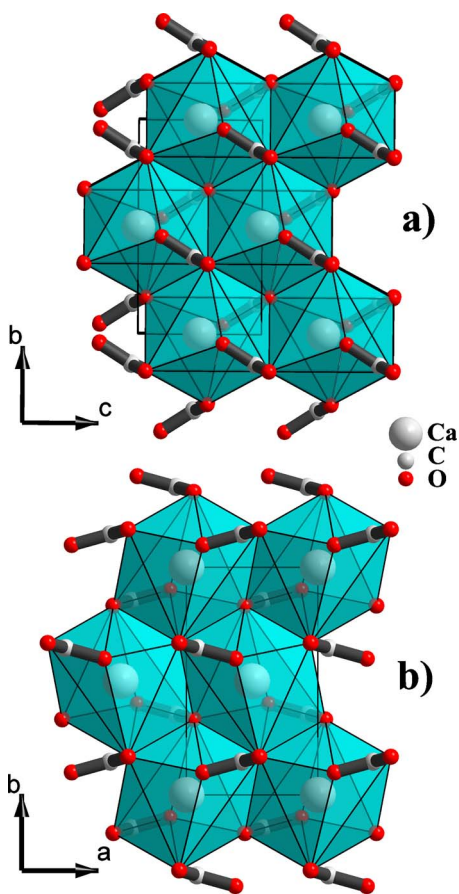


Figure 5. (Color online) Comparison of the *Ama2* model (a) with the *Pbnm* Meyer model (b). Projection along the longer axis (8.47 Å) showing a layer of edge-sharing  $\text{CaO}_8$  cubes as can be found by selecting a part of the  $\text{CaF}_2$  fluorite structure.

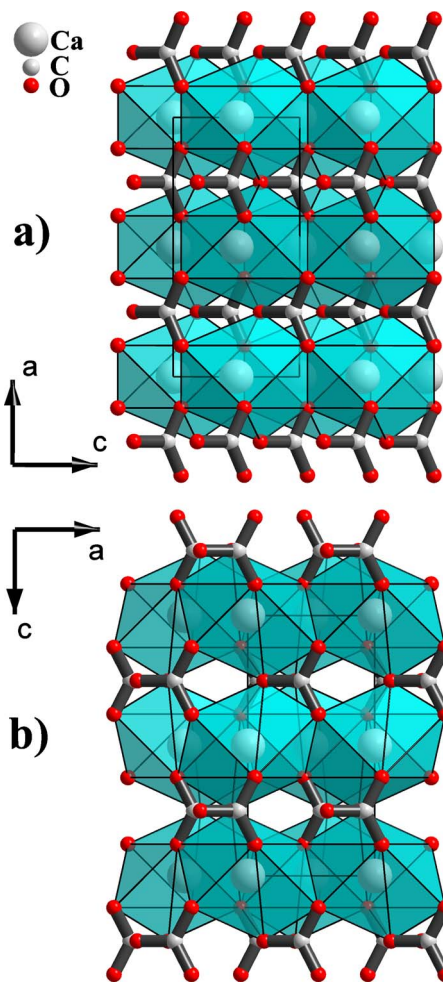


Figure 6. (Color online) Comparison of the *Ama2* model (a) with the *Pbnm* Meyer model (b). Projection along the *b* axis (7.16 Å) showing the differences in orientation and position of the  $\text{CO}_3$  groups in a similar matrix of  $\text{CaO}_8$  cubes interconnected by edges and corners.

morphs correspond to corner and full edge sharing of the  $\text{CO}_3$  “triangles.” In aragonite, this is operated through a slight but significant displacement of the carbon atom away from the oxygen plane (around 0.05 Å). In the newly proposed vaterite subcell, the lack of edge sharing of  $\text{CO}_3$  groups is compensated by the stabilization of  $\text{CaO}_8$  cubes, at the expense of around 2 kJ/mol compared to aragonite. Therefore, the development of vaterite during the first crystallization stages, when surface energy predominates, appears more understandable. The small number of additional weak superstructure reflections indicates that a subtle additional long range ordering modifies the average structure described in the present subcell. Seeing clearly only five peaks when the total number of reflections up to  $120^\circ 2\theta$  is supposed to increase from 51 to 246 from the hexagonal subcell to the supercell (and more in orthorhombic symmetry), it looks that the task of fully characterizing the supercell will remain uneasy and uncertain for a while. Moreover, an intense microtwinning may involve a non-negligible percentage of the total number of atoms (Figure 7) and be the reason for additional streaks. Faults and microtwinning are favoured by the fluorite-type structure edge-sharing  $\text{CaO}_8$  cubes being able to develop the observed layers in several space directions, accommodating easily, just like in  $\text{CaF}_2$ . Moreover, on

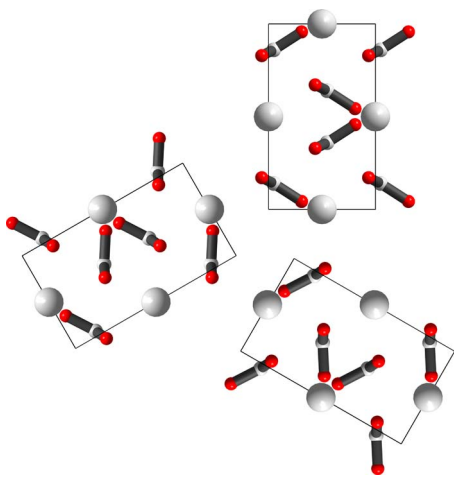


Figure 7. (Color online) The twinning hypothesis with three orthorhombic domains (space group *Ama2*) rotated by  $120^\circ$  along the *a* axis, regenerating the apparent hexagonal symmetry on a “single crystal.” This implies no change on the Ca atoms, only small oxygen atoms moves, and mainly C atoms positioned at the center of different  $O_3$  triangles generated by the edge sharing of the  $CaO_8$  cubes. The twin plane may well be orthogonal to the *a* orthorhombic axis (*c* in the old hexagonal description) rather than parallel to it.

one side or the other relative to the  $CaO_8$  edge-sharing cube layers, the  $CO_3$  group order may easily differ because the energy barrier is very small. The stacking may thus show the three orientations at  $120^\circ$  on the same pseudo-single-crystal along the *a* orthorhombic axis in the *Ama2* cell (*c* axis in the old hexagonal description) as observed by Qiao and Feng (2007). Our vaterite particles are globally  $4 \mu\text{m}$  diameter spheres, with SEM-visible substructure as small as acicular-like and radially disposed crystals of typically 50 nm diameter (Ouhenia *et al.*, 2008b) and a mean isotropic coherent domain size of 16 nm. Our microtwinning hypothesis then takes place in the formation of these small crystals rather than in an eventual formation of hexagonal-like vaterite platelets as sometimes observed in the literature and is intrinsically linked to the relatively higher energy state, easily transformed into calcite and/or aragonite.

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